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Fuzzy petrology in the origin of carbonatitic/pseudocarbonatitic Ca-rich ultrabasic magma at Polino (central Italy)

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The small upper Pleistocene diatreme of Polino (central Italy) is known in literature as one of the few monticellite alvikites (volcanic Ca-carbonatite) worldwide. This outcrop belongs to the Umbria-Latium Ultra-alkaline District (ULUD), an area characterized by scattered and small-volume strongly SiO₂-undersaturated ultrabasic igneous rocks located in the axial sector of the Apennine Mts. in central Italy. Petrographic and mineralogical evidences indicate that Polino olivine and phlogopite are liquidus phases rather than mantle xenocrysts as instead reported in literature. The presence of monticellite as rim of olivine phenocrysts and as groundmass phase indicates its late appearance in magma chambers at shallow depths, as demonstrated by experimental studies too. The absence of plagioclase and clinopyroxene along with the extremely MgO-rich composition of olivine (Fo₉₂₋₉₄) and phlogopite (average Mg# ~93) suggest for Polino magmas an origin from a carbonated H₂O-bearing mantle source at depths at least of 90–100 km, in the magnesite stability field. In contrast with what reported in literature, the ultimate strongly ultrabasic Ca-rich whole-rock composition (~15–25 wt% SiO₂, ~31–40 wt% CaO) and the abundant modal groundmass calcite are not pristine features of Polino magma. We propose that the observed mineral assemblage and whole-rock compositions result mostly from the assimilation of limestones by an ultrabasic melt at a depth of ~5 km. A reaction involving liquidus olivine + limestone producing monticellite + CO₂ vapour + calcite is at the base of the origin of the Polino pseudocarbonatitic igneous rocks.

The axial portion of the Apennine fold-and-thrust belt in central Italy is characterized by the sporadic presence of upper Pleistocene low-volume alkali- and lime-rich igneous rocks belonging to the kamafugite-melilitite-carbonatite clan¹⁻⁴, erupted along fault systems bordering Plio-Pleistocene grabens. Two of these outcrops near San Venanzo and Cupaello villages are type localities of strongly SiO₂-undersaturated kalsilite-bearing rocks named venanzite (a melanocratic variety of leucite olivine kalsilite melilitite) and coppaelite (a phlogopite kalsilite pyroxene melilitite without olivine)⁵. A ~10–15 m wide diatreme of Ca-rich rocks crops out in the Polino village in between San Venanzo and Cupaello (Fig. 1). These rocks, dated with ⁴⁰Ar/³⁹Ar ~246 ka⁶, so far have been classified as monticellite alvikites, i.e., volcanic Ca-carbonatites⁷⁻¹⁰.

The identification of Ca-carbonatitic volcanic activity has been considered a key feature to unravel the geodynamic setting of the Plio-Quaternary volcanism of central Italy. In particular, the presence of carbonatitic magmas has been proposed to deny the existence of active or fossil subduction tectonic settings along peninsular Italy, proposing the existence of a deep seated variably asymmetric (eastward-pushing) mantle plume or hot spot¹¹⁻¹³. From a general point of view, “a strong evidence for the link between the carbonatite genesis and the locations of deep-mantle plumes” has been recently proposed¹⁴.

The mantle source of the calcite-, melilitite- and kalsilite-bearing rocks is assumed to be the same of the potassic and ultrapotassic volcanic rocks of central-southern Italy (Latium-Campania administrative regions¹¹). On this

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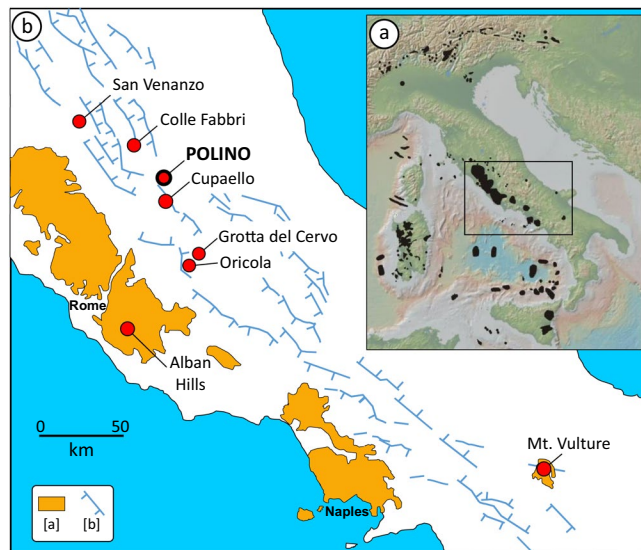


Figure 1. (a) Occurrence of the main subaerial and submerged Cenozoic igneous districts of circum-Tyrrhenian and Alpine realms. (b) Simplified geological map showing the main outcrops of carbonatites of Italy reported in literature. Red circles indicate ULUD outcrops, carbonatites and strongly SiO₂-undersaturated ultrapotassic volcanic rocks some of which as root-less pyroclastic facies. Alban Hills and Mt. Vulture are volcanoes where carbonatites or magmatic calcite have been reported. a = Plio-Quaternary potassic to ultrapotassic volcanic rocks; b = normal faults.

ground, any subduction-related process has been considered highly improbable for the entire Plio-Quaternary volcanism of peninsular Italy too (e.g., refs^{13,15}).

On the other hand, the majority of researchers considers the potassic and ultrapotassic volcanic rocks of peninsular Italy (including the rare kamafugitic-melilititic sub-groups) as occurring in a subduction-related setting, interpreting the lime-rich compositions like hybrid melts or paralavas (e.g., refs^{4,16–21}). The Pleistocene carbonatitic lava flow (alvikite) found at Mt. Vulture are interpreted as generated from liquid unmixing at crustal levels starting from nephelinitic to melilititic parental magmas, these latter generated above a slab window in a metasomatized mantle wedge²².

The aim of this work was to investigate the origin of Polino volcanic rocks (Terni, Italy), so far classified as Ca-carbonatite because of their abundance of calcite and overall Ca-rich and Si-poor whole-rock composition. The coexistence of Mg-rich olivine, Mg-rich phlogopite, monticellite, globular and sparry calcite and the absence of feldspars and pyroxenes make Polino rocks unique worldwide, distinct from the other rare cases of monticellite carbonatites reported in literature.

On the basis of petrographic, mineral chemical and whole-rock geochemical evidences, as well as petrologic considerations, we propose an alternative petrogenetic process. Our main conclusions can be summarized as: (1) olivine and phlogopite macrocrysts are liquidus phases rather than mantle debris; (2) the absence of plagioclase and pyroxene along with the high Mg# of olivine and phlogopite indicate strongly SiO₂-undersaturated magma compositions (CIPW larnite normative) formed by melting of a magnesite-bearing peridotitic mantle; (3) the actual mineral assemblage of Polino rocks consisting of olivine, phlogopite, monticellite and calcite can be explained in light of shallow-depth digestion of country rocks represented by a ~5 km-thick limestone succession; (4) the Polino igneous rocks are better classified as pseudocarbonatites or, more properly, endoskarns; (5) carbonatites, commonly considered in literature to be indicative of within-plate or continental rifting settings, connected to the presence of mantle plumes^{13,14,23,24}, cannot be used to infer geodynamic environments.

Carbonatites

Carbonatites are igneous rocks composed of >50% modal primary carbonate and with <20 wt% SiO₂ (ref.⁵). Carbon must be juvenile – magma-derived – and not from wall rocks or external hydrothermal fluids. Calcite-carbonatite (also defined as calcio-carbonatite) magmas are by far the most abundant lithologies of this extremely rare rock group. Carbonatites can be either alkali-rich or alkali-poor, and are commonly associated with silicate rocks characterized by alkaline ultrabasic compositions such as nephelinite/nephelinolite, melilitite/melilitolite, essexite, aillikite, melteigite and kimberlite, but can also be genetically related to evolved and SiO₂-richer lithologies such as trachyte/syenite, phonolite/nepheline syenite²⁵. Albeit rare²⁶, carbonatitic inclusions have been reported in mantle xenoliths found in mafic sodic magmas (e.g., refs^{27,28}). On the other hand, mantle xenoliths have been never²⁹ or rarely³⁰ recorded in carbonatitic magmas because of their preferential settling in low viscosity and low density magmas.

The origin, the evolution and the plate tectonics significance of carbonatites have been debated for decades with several models proposed over the last 50 years. Mitchell (ref.³¹) described three main types of carbonatites: (1) Primary Carbonatites – genetically related to mantle-derived magmas such as nephelinite, melilitite and

kimberlite. These magmas form by partial melting of a carbonated peridotite or eclogite (e.g., ref.³²), likely followed by fractional crystallization (e.g., ref.³³), and liquid immiscibility from alkali-rich to alkali-poor silicate magmas (e.g., refs^{29,34–36}); (2) Carbothermal Carbonatites³⁷ – related to low-temperature fluids derived from fractionated magmas enriched in CO₂, H₂O and fluorine; (3) Pseudocarbonatites – formed by anatectic melting of crustal rocks, also named calc-silicate skarns or anatectic calc-silicate veins. Barker (ref.²⁹) also considers the problem of pseudocarbonatites, defining them carbonate-rich rocks with hydrothermal, metamorphic or sedimentary origin, to be distinguished by true igneous carbonatites. Lentz (ref.³⁸) considers pseudocarbonatites in terms of skarn reaction equilibria too.

Gudfinnsson and Presnall (ref.³⁹) showed experimentally that the composition of melts generated by melting of a carbonated peridotite is strictly related to the temperature and depth of the source rock. With increasing temperature, between 2 and 8 GPa, different melts can be produced, with compositions ranging from carbonatitic (<10 wt% SiO₂; CaO >25 wt%) to melilititic (~20–30 wt% SiO₂; CaO ~20–25 wt%), kimberlitic (~30–40 wt% SiO₂; CaO <20 wt%) and eventually basaltic (~>40 wt% SiO₂; CaO <10 wt%). Below the solidus, the stable carbonate is dolomite at P ~2–3 GPa, and magnesite at P >3 GPa. Carbonatitic melts generated from a carbonated peridotite contain dolomitic to magnesian phases rather than calcite/aragonite as instead observed in case of melting of Ca-rich eclogite assemblages⁴⁰.

As concerns the circum-Mediterranean area, the presence of carbonatitic rocks (or carbonatitic components) have been reported in SE Libya (Uwaynat), NW Morocco (Taouirt, Tamazert, Saghro), central Spain (Calatrava), central France (Chabrières, Clermont Ferrand), Germany (Kaiserstuhl, Eifel), Bohemian Massif (Czech Republic, Flurbuhl) eastern and central Anatolia (Malatya, Sivas, Kizilcaoren) and in central and southern Italy (Polino, Alban Hills, Mt. Vulture; ref.⁴¹ and references therein).

One of the major problems in identifying true carbonatitic rocks is to distinguish magmatic textures from later replacement and secondary textures (e.g., ref.⁴²). Indeed, subsolidus plastic flow, deuteric alteration and solution-precipitation recrystallization of magmatic calcite can lead to a difficult distinction from hydrothermal and metasedimentary carbonate rocks²⁹. Also the presence of minerals typical of carbonatite rocks cannot be considered a proof for their mantle origin, being these phases also found in calc-skarns^{38,43}. We report petrographic and mineralogical evidences according which the carbonate-rich fraction of Polino rocks cannot be considered a feature of primary – i.e., in equilibrium with mantle rocks – conditions.

Carbonatites in Italy

Rocks interpreted as carbonatites crop out in central and southern Italy (Fig. 1) as carbonatites and carbonatite-related lithologies (Polino, Cupaello, Colle Fabbri, Oricola and San Venanzo) of the *Umbria-Latium Ultra-alkaline District* (ULUD; ~0.64–0.26 Ma) also known as *Intra-Apenine Province* (IAP), *Intra-Apenine Volcanism* (IAV) or *Intramontane Ultra-alkaline Province* (IUP^{1,4,6,11,44–47}). These rocks are sometimes associated with strongly SiO₂-undersaturated melts such as melilitites and kamafugites, as well as with minor evolved compositions (trachyphonolites); only in the Polino case no associated silicatic magma is present being the monticellite alvikite the only reported igneous rock. The typical (but not ubiquitous) mineral paragenesis of ULUD rocks is represented by variable amounts of foids (mostly leucite and nepheline-kalsilite solid solution), melilite, phlogopite, olivine and clinopyroxene.

Southern Italy carbonatites are reported in Mt. Vulture volcanic complex as a small alvikitic lava flow²², as plutonic xenoliths (sövite⁴⁸) and as carbonate inclusions in mantle xenoliths⁴⁹. High δ¹⁸O clinopyroxenes (up to +8.3‰), the correlation between δ¹⁸O and Si-Al^{IV} content in clinopyroxene, Mg#-rich clinopyroxene (up to 95) and phlogopite (up to 93), Fo-rich olivine (up to 98), abundant skarn xenoliths, high-energy explosive eruptions (e.g., Villa Senni) and groundmass calcite in Alban Hills lavas and granular ejecta are interpreted as a result of assimilation of limestone wall rocks by basic primary melts (e.g., refs^{43,50–55}). This is modelled as open system skarnification of a basic silicatic melt which digested limestone walls of a shallow magma chamber, evolved via fractional crystallization towards the silicate-carbonatite solvus, generating carbonatitic melts with clear secondary origin⁵⁴.

Carbonatites from central Italy are associated to small monogenetic centres (diatreme structures) pyroclastic rocks and minor lavas with ultrapotassic composition located close to the Apennine Chain axis (Fig. 1). Overall, the trace element content and Sr-Nd-Pb isotopic ratios of Italian carbonatites match those of the Roman Province rocks⁴⁷, but their origin remains highly debated. Stoppa *et al.* (ref.⁵⁶) interpret the carbonate-rich lithologies associated with the ultrapotassic rocks as true Ca-carbonatitic melts, while Peccerillo (ref.⁵⁷) relates them to interaction between silicate melt and sedimentary limestone country rocks.

According to several authors, the volcanic rocks of Polino, classified as monticellite calcio-carbonatites^{7–10,45,58}, are characterized by abundant mantle xenoliths and xenocrysts (forsterite and phlogopite). Here, we demonstrate that the petrographic features, the mineralogy and the chemical composition of Polino ultrabasic rocks are more compatible with shallow depth assimilation of sedimentary carbonate wall rocks by SiO₂-poor Mg-rich carbonated melts.

Polino

Despite the limited outcrop (<100 m²), the Polino volcanic rocks occur in two facies. The first (pyroclastic facies) is characterized by the presence of mm- to cm-sized droplets of silicate melt agglutinated by sparry calcite cement, while the second show a more massive aspect (Fig. 2). Both facies are characterized by abundant euhedral to subhedral forsteritic olivine (Fo_{92–94}) followed by euhedral phlogopite (Mg# 90–94) and monticellite, coexisting with accessory phases as Ca-Ti/Ca-Si perovskite, schorlomite garnet and Fe-Ti oxides, all associated with variable amounts of sparry, sedimentary or amoeboid calcite (e.g., ref.⁵⁹; Fig. 2). All the minerals are SiO₂-poor to SiO₂-free phases, i.e., to complete occupancy of tetrahedral sites in the schorlomite garnet, all the Al and half of Fe must be assigned to tetrahedral coordination⁵⁹. These rocks are characterized by the absence of feldspars and

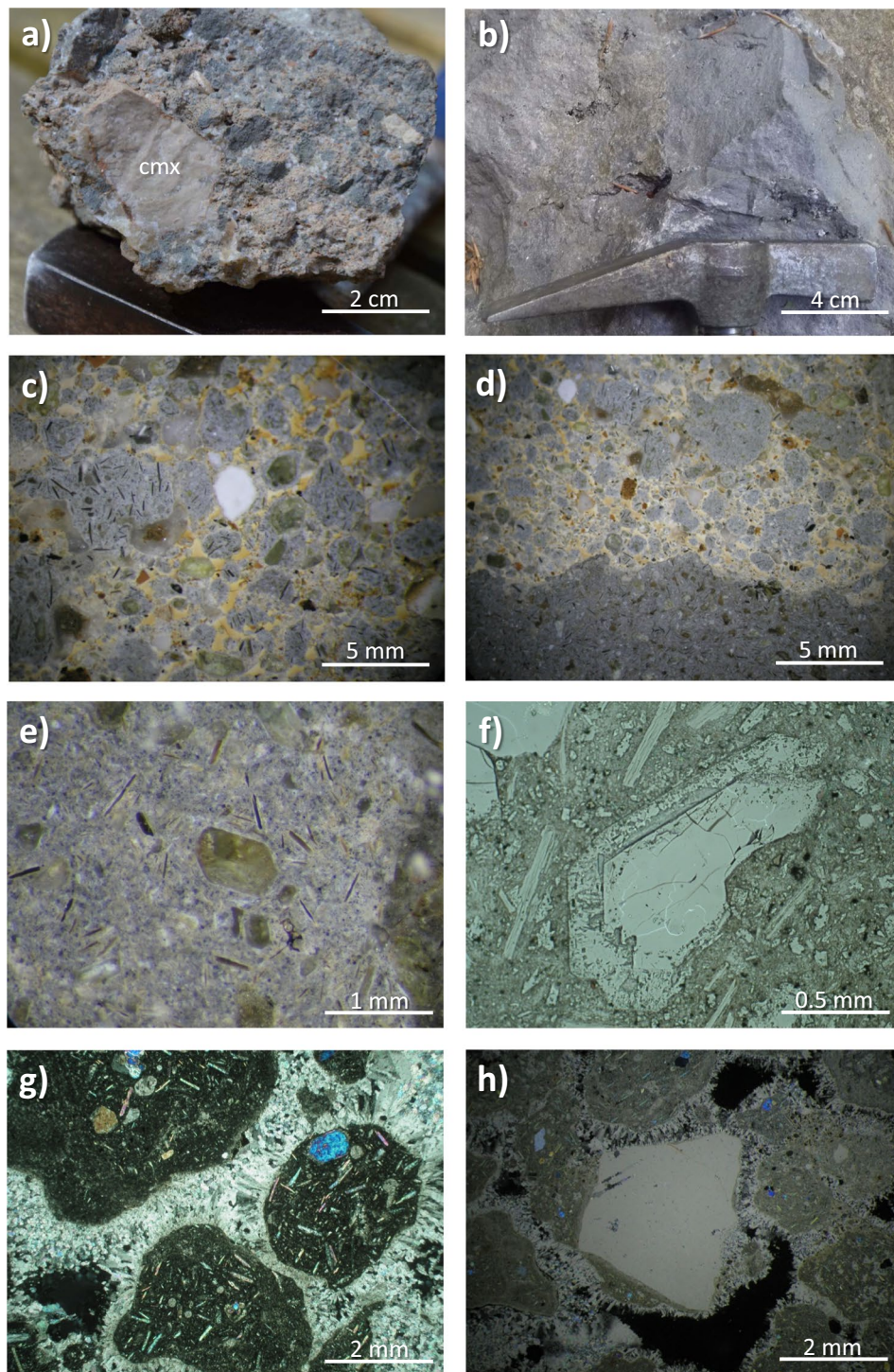


Figure 2. Polino volcanic rocks. (a) Pyroclastic facies with a large Calcare Massiccio limestone xenolith (cmx), smaller melt droplets (greyish) and smaller limestone xenoliths agglutinated in an havana colour calcitic cement. (b) Massive facies. (c) Close-up of polished pyroclastic facies with silicatic melt droplets (greyish) agglutinated by sparry calcitic cement (yellowish to brownish) and Calcare Massiccio limestone xenoliths (white) with olivine (greenish) and phlogopite (acicular habit) phenocrysts. (d) Close up of polished contact between pyroclastic (top) and massive (bottom) facies. Only the massive facies samples have been selected for whole-rock analyses. (e) Close up of polished massive facies with euhedral olivine and phlogopite phenocrysts. (f) Plane-polarized view of euhedral olivine phenocryst rimmed by monticellite rim, set, together with phlogopite phenocrysts and micro-phenocrysts, in a fine-grained monticellite- and calcite-rich groundmass. (g) Crossed polar view showing silicatic globules – characterized by high-birefringent olivine and phlogopite phenocrysts and havana calcite globules – cemented by sparry radiating calcite. (h) Crossed polar view showing a large Calcare Massiccio limestone xenolith in a scoriaceous (black = voids) pyroclastic facies with silicatic globules cemented by sparry radiating calcite.

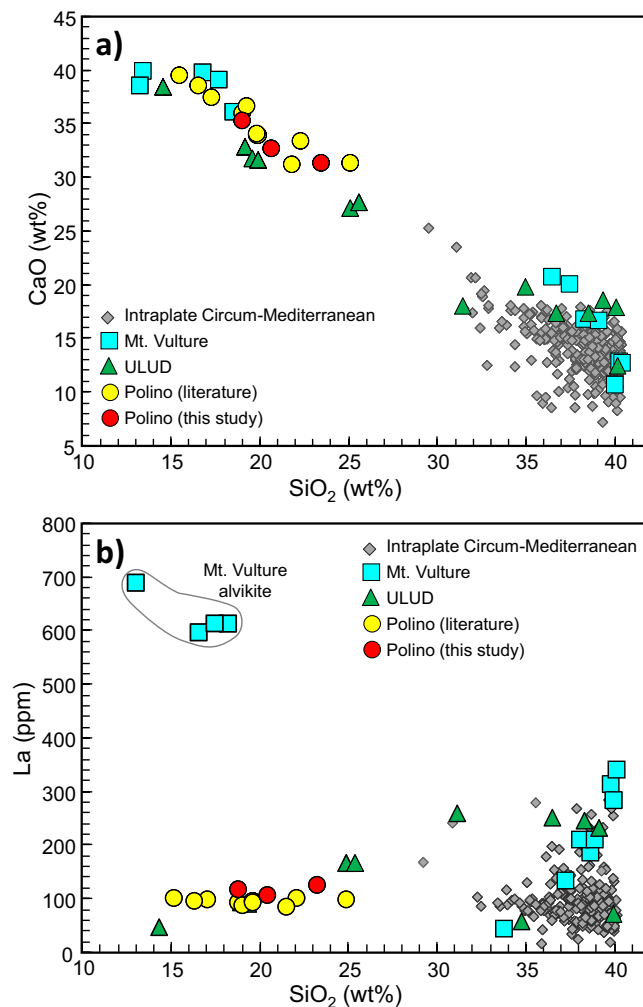


Figure 3. CaO vs. SiO₂ (a) and La vs. SiO₂ (b) whole-rock composition of Polino volcanic rocks. The composition of other ULUD (Umbria-Latium Ultra-alkaline District) rocks, as well as Mt. Vulture and circum-Mediterranean igneous rocks with SiO₂ < 40 wt% are reported for comparison (ref.²¹ and references therein).

pyroxenes, with the exception of very rare (<0.1% vol) occurrences of sanidine and diopside xenocrysts from older phonolitic pyroclastic cover pierced by the diatreme⁸. The Polino diatreme crosses the Calcare Massiccio Formation, a Lower Jurassic peritidal carbonate succession cropping out in the Apennines originally deposited along the passive margin of the Adria micro-plate (e.g., ref.⁶⁰). During two campaigns in 2016 and 2017, twelve samples were collected and analysed for petrographic investigation using the polarizing microscope, scanning electron microscopy (SEM) and electron microprobe (EMP), while three of them have been analyzed for major oxides (ICP-AES) and trace elements (ICP-MS) at the Activation Laboratories (Ontario, Canada). Only massive facies samples have been chosen for analyses, excluding pyroclastic facies rocks rich in sparry calcite cementing silicate globules. About 0.2 g sample have been thermally decomposed in a resistance furnace in a pure nitrogen environment at 1000 °C, directly releasing CO₂. H₂O is removed in a moisture trap prior to the detection of CO₂ in the infra-red cell. Full details in <http://www.actlabs.com>. The EMP details are reported in Lustrino *et al.* (ref.⁶¹).

Some of the major oxides of the three Polino volcanic rocks analyzed in this study (full dataset in Table A1) are reported in Figs 3 and 4, together with the Polino literature data, the composition of the ULUD, Mt. Vulture rocks and other intraplate circum-Mediterranean ultrabasic rocks with SiO₂ content < 40 wt%. The LOI of the Polino rocks is very high (19.8–25.2 wt%), as expected for calcite-rich compositions.

The rocks analyzed in this study show low SiO₂ contents (~18.8–23.8 wt%), within the composition reported in literature of ~15.1–24.8 wt% (refs^{7,8,45}). The range in silica content is surprisingly wide considering the outcrop extension (<100 m²) and is here explained as due to the difficulty to completely separate the sparry cement from the silicatic portion in the pyroclastic facies rocks. The CaO content ranges from 31.6 to 39.7 wt% and shows negative correlation with SiO₂ (R² = 0.89). With the exception of two literature analyses, CaO is correlated with CO₂ (14.5–26.0 wt%; R² = 0.81; Fig. 4). These correlations highlight two components, one relatively SiO₂-rich but CaO- and CO₂-poor, and the other SiO₂-poor (to SiO₂-free) and Ca-CO₂-rich. The second term has been considered the most representative Italian carbonatite end-member by Martin *et al.* (ref.¹⁰) and is reported in several worldwide carbonatite lists (e.g., refs^{14,29,30,62–64}). Polino rocks are alkali-poor, with Na₂O < 0.15 wt% and K₂O

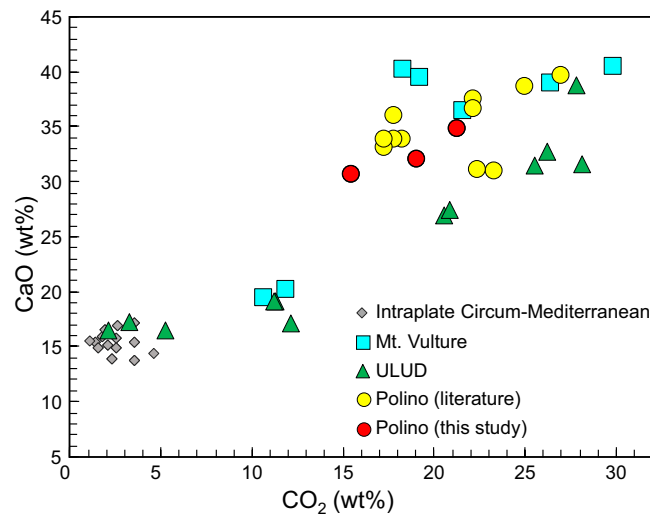


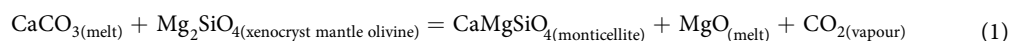
Figure 4. CaO vs. CO₂ whole-rock composition of Polino volcanic rocks. The composition of other ULUD (Umbria-Latium Ultra-alkaline District) rocks, as well as Mt. Vulture and circum-Mediterranean igneous rocks with SiO₂ < 40 wt% are reported for comparison (ref.²¹ and references therein).

<0.92 wt%, a feature in contrast with the coeval and nearby potassic to ultrapotassic composition of the Italian Pleistocene volcanic rocks and the other ULUD products^{4,20,21,42}.

Compared with the average Ca-carbonatite compositions^{65,66}, the average composition of Polino rocks shows much lower Sr (1690 ppm vs. 7272 ppm), Nb (15 ppm vs. 343 ppm), Ta (0.7 ppm vs. 9.1 ppm) and LREE (e.g., Ce = 217 ppm vs. 1687 ppm), coupled with much higher Rb (63 ppm vs. 14 ppm), Cr (526 ppm vs. 13 ppm) and Ni (320 ppm vs. 18 ppm). A comparison with natural monticellite carbonatites shows the peculiarity of Polino rocks. Indeed, Polino rocks, compared with Fort Portal (Uganda) monticellite Ca-carbonatite lavas and tuffs, have lower TiO₂ (<0.6 wt% vs. >1.6 wt%), Fe₂O_{3tot} (<5.3 wt% vs. >9.0 wt%), Na₂O (<0.19 wt% vs. >0.37 wt%), P₂O₅ (<0.6 wt% vs. >1.6 wt%), V (<55 ppm vs. >207 ppm), Sr (<1897 ppm vs. >3324 ppm), La (<122 ppm vs. >250 ppm), Nb (<21 ppm vs. >241 ppm) and Ta (<0.8 ppm vs. >10.6 ppm), as well as higher Cr (>368 ppm vs. <142 ppm) and Ni (>231 ppm vs. <73 ppm). Moreover, Fort Portal carbonatitic lavas and tuffs are characterized by the association of Mg-rich calcite (MgO up to 13 wt%) and calcite, while Polino rocks host Mg-poor calcites (MgO <0.6 wt%) only. Fort Portal carbonatites are also associated to abundant silicatic magma (mostly melilitites), while in Polino no associated silicate rocks are present (being the silicatic component of Polino interpreted in literature as related to the contamination of a pure Ca-carbonatitic melt with matrix peridotite *en route* to the surface).

Whole Rock and Mineral Chemical Constraints for the Origin of Polino Rocks

The relationship between olivine and monticellite is the key aspect to unravel the origin of the Polino magma. Polino monticellite has been interpreted as the result of the interaction between a Ca-carbonatitic mantle melt with xenocrystic mantle forsterite. According to several authors (refs^{7-9,58}), the Ca-carbonatitic melt should be generated at mantle depths by immiscibility from a carbonatite-melilitite liquid. Then, it partially reacted with portions of mantle wall rocks during its ascent to the surface, with the chemical interaction between liquid carbonatite and mantle olivine being expressed in the form⁹:



This reaction would result in a dilution of the original CaO content of the carbonatitic magma coupled with an increase of its SiO₂ and MgO content. According to Bell and Kjarsgaard (ref.⁶⁷; p. 86) the presence of mantle xenoliths at San Venanzo and Polino “is a testament to the rapid ascent of the kamafugitic and carbonatitic magmas”, rendering “highly problematic” an open system interaction between a silicatic (kamafugitic) magma and local limestone. On the other hand, Barker (ref.²⁹; p. 46) underlines that ultramafic mantle xenoliths have never been reported in carbonatite lavas or intrusions, due to “the low viscosity and low density of carbonatitic liquid or to the separation of an immiscible liquid from mantle-derived silicate magma within a crustal reservoir. In both cases, ultramafic xenoliths would settle out”.

Fort Portal carbonatite contains a variable cargo of silicatic rocks as gneissic, amphibolitic, and gabbroic xenoliths. In order to calculate the relatively pristine composition of the carbonatitic magma, Eby *et al.* (ref.⁶⁸), following the approach of Barker and Nixon (ref.⁶⁹), subtracted the silicatic component by regressing major and trace elements vs. alumina to Al₂O₃ = 0. In many cases (SiO₂, TiO₂, MnO, P₂O₅, Cr, Ni, Zn, REE, Y, Nb, Zr, Ta, Th) the R² value obtained with these regressions is very high (>0.9), indicating the petrological sound of such a procedure. If olivine and phlogopite in Polino rocks are mantle xenocrysts, we should obtain similar results. However, a regression of both major and trace elements vs. Al₂O₃ do not show any appreciable correlation, with major elements showing R² values <0.39 and trace elements showing R² values <0.33.

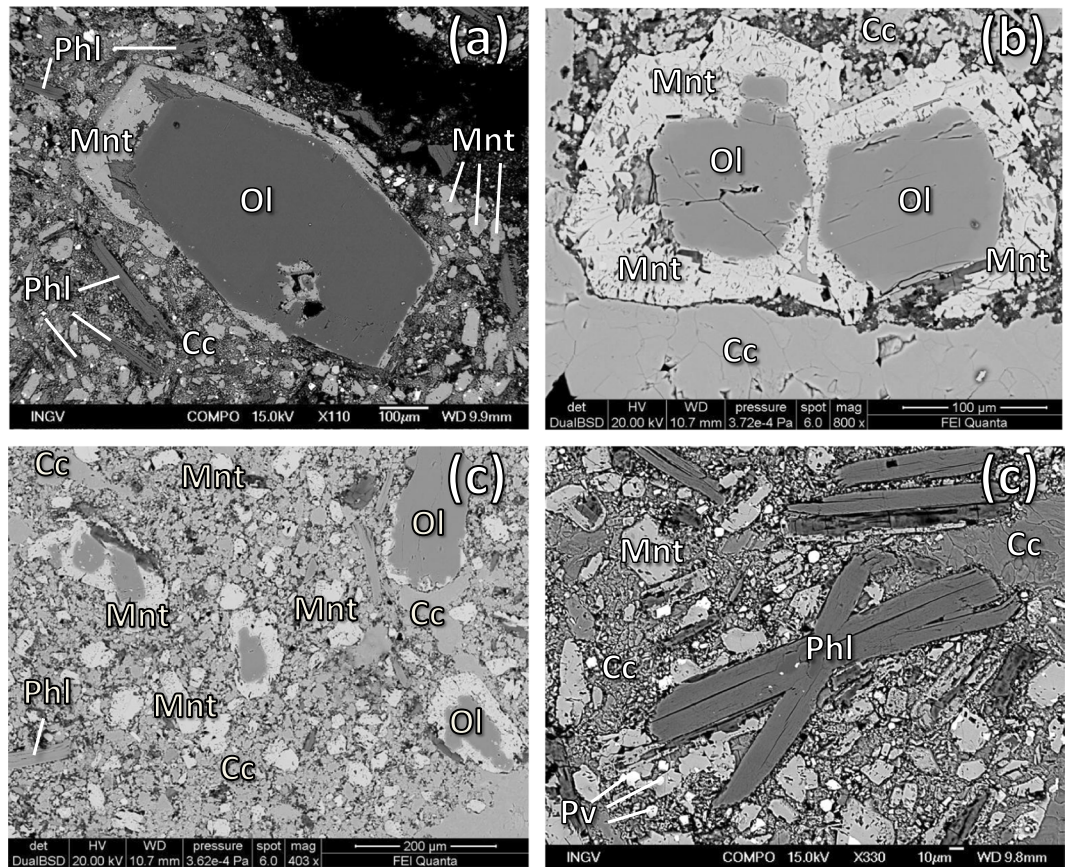


Figure 5. (a) SEM images of Polino euhedral olivine (Ol) rimmed by monticellite (Mnt), groundmass monticellite, euhedral phlogopite (Phl) and ameboid calcite (Cc). (b) SEM image of Polino euhedral olivine rimmed by monticellite close to a border of sparry calcite. (c) SEM image of Polino rock petrography. (d) SEM image of Polino euhedral phlogopite laths in a groundmass made up of monticellite, calcite and perovskite (Pv).

A microstructural and compositional study of the main mineral phases found in Polino rocks allows us to exclude a xenocrystic origin for forsterite and phlogopite. These two minerals, indeed, are characterized by euhedral to subhedral shape, with phlogopite mostly represented by tiny elongated euhedral laths (Figs 2 and 5; see also pictures in Table A2). Interestingly, the age of Polino rocks (~246 ka) has been estimated via $^{40}\text{Ar}/^{39}\text{Ar}$ on phlogopite separates, indicating that this is a phase crystallized from a melt rather than a mantle xenocryst.

The euhedral shape strongly differs from olivine crystals found in mantle xenoliths (but also in Alpine-type ophiolitic massifs or dredged peridotites), always characterized by anhedral habit. The interpretation of Polino olivines as mantle xenocrysts seems, therefore, highly improbable (e.g., ref.⁷⁰). In addition, the absence of any deformation texture in the olivine crystals in Polino olivines is a further proof at odd with a mantle xenocrystic origin hypothesis, supporting the derivation as a liquidus phase. The lack of peridotite mantle xenoliths or mantle xenocrysts would be sufficient to confute the reaction (1).

Our new interpretation of Polino olivines as liquidus phase is also supported by EMP data. The forsterite content of olivine (expressed as Fo-Fa solid solution) is particularly high, mostly clustering in the 91.8–94.0 range (463 analyses; average = 93.2, st. dev. ± 0.34 ; Table A2), much higher than the Mg# of monticellite (75.8–85.1; 83 analyses; average 79.2, st. dev. ± 2.3 ; Table A2). This high forsterite content has been considered the strongest proof for a mantle xenocrystic origin (e.g., refs^{7–9,58}). In the CaO vs. Fo diagram (Fig. 6), the Polino olivines are compared with worldwide mantle xenolith olivines from the PETDB database (<https://www.earthchem.org/petdb>). Polino olivines are distinct, being characterized by much higher CaO content (0.16–1.9 wt%) than mantle olivine (<0.12 wt% CaO for >95% of the data). The major element content of Polino olivines also reflects their liquidus origin, albeit the anomalously high Fo content requires specific interpretation (see below). Similarly, phlogopite is characterized by Mg-rich compositions with Mg# ratios ($\text{Mg}\# = 100 * [\text{Mg}/(\text{Mg} + \text{Fe}^{2+})]$) ranging from 90.1 to 94.5 (59 analyses; average 93.19, st. dev. ± 0.39 ; Fig. 7a and Table A3), but plot far from the field of typical mantle phlogopites when minor elements are taken into account (e.g., Na₂O vs. CaO; Fig. 7b).

Polino olivines are characterized by relatively high Ni (ranging from ~550 to ~5700 ppm; average = ~3300 ppm, st. dev. ± 1180 ppm; Table A2). According to Ammannati *et al.* (ref.⁷¹), the Italian plagioclite rocks have olivine poor in Ni (~1100–1960 ppm; average ~1330 ppm) because generated from a peridotitic mantle metasomatized by CaCO₃-rich agents. On the other hand, SiO₂-richer lamproitic rocks are characterized by much higher Ni (~1960–4950 ppm; average ~3140 ppm). The reason of Ni enrichment in olivine from lamproitic melts would be the derivation of these liquids from olivine-poor to olivine-free mantle source. The

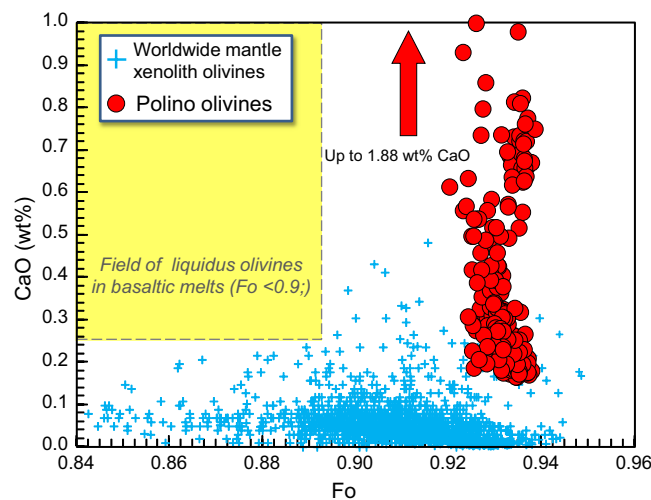


Figure 6. CaO vs. Fo diagram for Polino olivines. The analyses of worldwide mantle xenolith olivines taken from PEDTB database are reported for comparison.

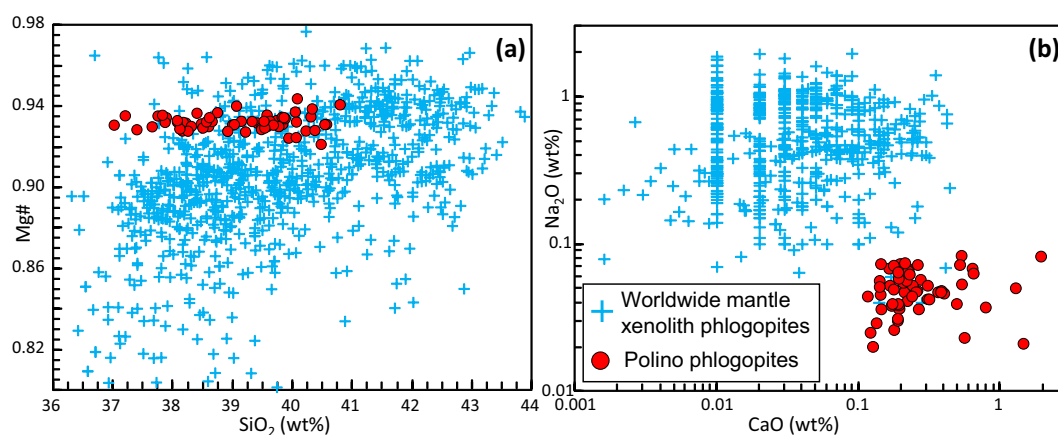


Figure 7. Polino phlogopite EMP data. Worldwide phlogopites in mantle xenoliths from GEOROC database, selecting rock types falling in dunite, harzburgite, lherzolite and peridotite definitions and with the following additional filters: $\text{TiO}_2 < 10 \text{ wt\%}$, $\text{Al}_2\text{O}_3 > 9 \text{ wt\%}$, $\text{Cr}_2\text{O}_3 < 3 \text{ wt\%}$; $\text{MgO} > 10 \text{ wt\%}$ and $< 30 \text{ wt\%}$, $\text{Na}_2\text{O} < 2 \text{ wt\%}$, $\text{Mg\#} > 0.80$). The complete list can be requested to the first author.

depletion in olivine would result after the reaction of mantle forsterite with silica-rich agents released from recycled terrigenous sediments, forming enstatite-rich metasomes^{72,73}. On the other hand, the low-Ni (and high-Ca; $\sim 2400\text{--}4000 \text{ ppm}$) content of olivines in the plagio-leucitites would reflect “the effect of the reaction between melts of carbonate-rich sediments with peridotite, stabilizing newly formed olivine and clinopyroxene at the expense of orthopyroxene” (ref.⁷¹; p. 72). Polino olivines do not have the low Ni content of the melts generated from a carbonate-metasomatized sources, but share with them the relatively Ca-rich ($\sim 430\text{--}6220 \text{ ppm}$; average $\sim 2720 \text{ ppm}$) content. This questions the validity of the general rule proposed by Ammannati *et al.* (ref.⁷¹).

The extremely MgO-rich composition of olivine up to Fo_{94} and phlogopite up to $\text{Mg\#} 94$, can be related either to: (1) a strongly depleted (i.e., Fe-poor) mantle source; (2) Fe-Mg distribution coefficient between liquidus olivine and coexisting primitive melt [$\text{Kd}(\text{Fe-Mg})_{\text{ol/melt}}$] much lower than what measured in melts produced from CO_2 -free peridotitic source with $\text{Kd}(\text{Fe-Mg})_{\text{ol/melt}}$ of $0.30 (\pm 0.03)$; ref.⁷⁴); (3) redox conditions resulting after CaCO_3 digestion, under which Fe^{2+} is oxidised to Fe^{3+} , hence limiting its incorporation into the olivine lattice (e.g., refs^{53,72}); (4) digestion of basement dolomitic rocks⁷⁵; (5) anomalously Mg-rich (e.g., dolomite/magnesite-bearing carbonated mantle) source.

Forsterite is never observed in the groundmass, but as phenocryst only, characterized by a variably thick monticellite rim. The chemical composition of Polino olivine is, also in the largest crystals, homogeneous (Table A2), suggesting that the melt from which the euhedral olivine crystallized was chemically homogeneous and did not change over time. An abrupt compositional variation occurs in the outer rims only, when CaO content sharply increases from values usually $< 0.8 \text{ wt\%}$ to $> 31 \text{ wt\%}$, typical of monticellite, coupled with MgO decrease from $> 50 \text{ wt\%}$ to $< 20 \text{ wt\%}$, only few microns across (Fig. 8; Table A2). The boundary between olivine and monticellite

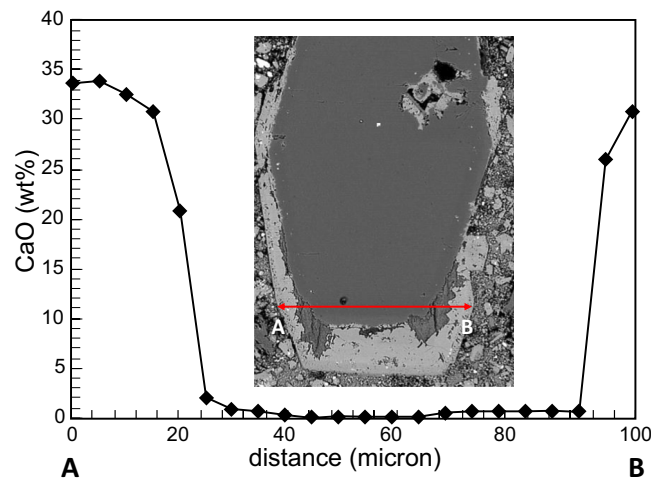


Figure 8. CaO transect of a compositionally homogeneous Polino olivine phenocryst rimmed by a thin monticellite rim seen in BSE and analyzed with EMP.

rim appears always abrupt and corroded, with signs of reaction and irregular contacts. This can be interpreted as a chemical shock in the form of a change of composition of the crystallizing melt.

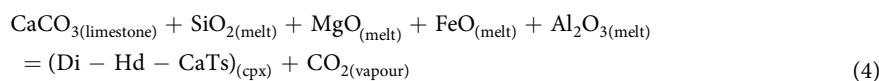
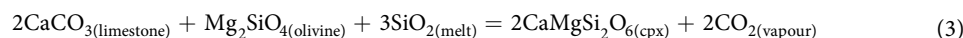
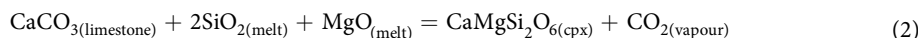
We interpret this feature as the indicator of the external input of CaO-rich lithologies in the silicatic magma when it started to pond at shallow crustal depths within sedimentary carbonate magma chamber at a pressure of ~200 MPa. During shallow depth ponding, the hydrous (because phlogopite-bearing) silicatic magma transferred latent heat of crystallization to the country rock, allowing partial dissolution of sedimentary calcite, strongly enriching the melt with CaO and CO₂ formed by calcite breakdown. Worth noting, the upwelling magma at Polino had to pierce a Jurassic dolostone and limestone succession ~5 km thick⁷⁶, and it is hard to believe that such a small volume of magma can have escaped significant interaction with wall rocks.

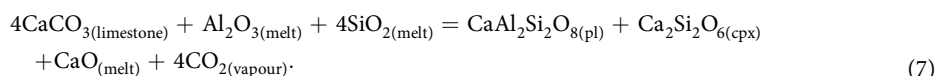
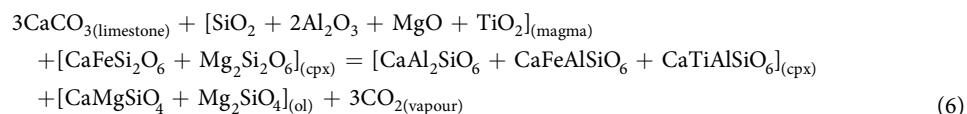
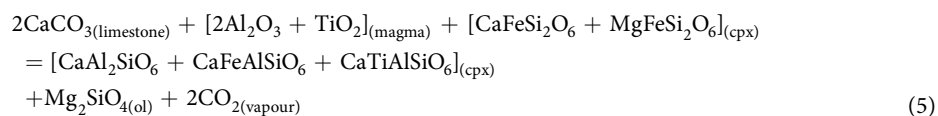
A peculiarity of Polino rocks is the common presence of phlogopite despite the low whole-rock K₂O content (average 0.63 wt% ± 0.21 wt%; Table A1). The high SiO₂ content of sanidine (~65 wt%) prevents its formation in strongly SiO₂-undersaturated melts such as those of Polino. Alternatively, anhydrous strongly ultrabasic K₂O-rich magma could stabilize kalsilite and/or leucite as the most stable K-bearing minerals. The absence of these two minerals is related to the overall low amount of K₂O in the magma, at odds with the high K₂O content of leucite (~22 wt%) and kalsilite (~30 wt%). The presence of dissolved H₂O and high MgO contents in the magma allowed the formation of phlogopite (~8.8–10.5 wt% K₂O) as the most stable K-bearing phase. The presence of this mica is, therefore, not to be considered as an anomalous feature, due to the specific conditions of the Mg-rich parental magma (before the interaction with limestones at shallow depths) and the availability of water in the magma.

Carbonate-Silicate Magma Interaction Styles

Based on textural and chemical evidences, we propose a two-stage process to explain the origin of olivine and monticellite. The first step includes the growth of compositionally homogeneous euhedral forsterite crystals as liquidus phases in equilibrium with an H₂O- and CO₂-bearing ultrabasic melt, followed by crystallization of phlogopite at water-saturated conditions. The second step is associated to an abrupt change of composition of the melt manifested by the appearance of monticellite at the expenses of forsterite around olivine phenocrysts and as groundmass phase not associated to any contemporary olivine crystallization. Once monticellite appeared, olivine stopped to grow. Groundmass monticellite is closely associated with precipitation of microcrystalline calcite after assimilation of limestone wall rocks.

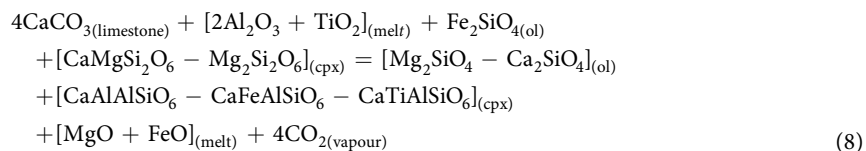
An important contribution over the last ten years has come from experimental studies on the melt composition after assimilation of carbonates. All the previous studies have used “basaltic” compositions (i.e., plagioclase-bearing melts) as starting materials to reproduce assimilation paths as function of temperature, pressure and *f*O₂. These experiments agree in the occurrence of olivine with relatively high Mg content (Mg# ~95) followed by the crystallization of clinopyroxene as result of the magma interaction with limestones according to the reactions:



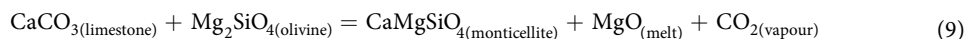


Equation (2) has been proposed by Iacono Marziano *et al.* (ref.⁷⁷) modelling natural basalts from Stromboli; Eq. (3) is from Iacono Marziano *et al.* (ref.⁷⁵) modelling the most primitive melt of Alban Hills (phono-tephrite); Eq. (4) is from Mollo *et al.* (ref.⁷⁸) modelling a synthetic glass resembling the most primitive K-basalt of the Roman Province; Eqs (5 and 6) are from and Mollo and Vona (ref.⁷⁹) modelling Mt. Vulsini shoshonite; Eq. (7) is from Carter and Dasgupta (ref.⁸⁰) modelling a composition similar to that used by Iacono Marziano *et al.* (ref.⁷⁵) and a Vesuvius shoshonite. Following these experimental evidences, some Alban Hills volcanic rocks have been interpreted as related to silicatic magma experiencing extensive carbonate assimilation (both limestone and dolostone) on the basis of the common presence of skarns and magmatic calcite^{51,54,77}.

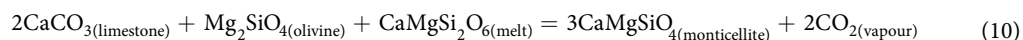
Other reactions include the increase of larnite component in olivine, but not the formation of pure monticellite⁸¹:



Also in this case, the major impact of carbonate (calcite/limestone) digestion on the mineralogy of hybrid magma is the increase of modal abundance of clinopyroxene and its increase in kushiroite ($\text{CaAl}_2\text{SiO}_6$), esseneite ($\text{CaFe}^{3+}\text{AlSiO}_6$) and grossmanite (CaTiAlSiO_6) components in clinopyroxene. In the case of Polino, the absence of clinopyroxene and plagioclase, as well as the presence of a monticellite rim around liquidus Mg-Fe olivine and as groundmass phase indicate strongly SiO_2 undersaturated compositions of the pristine (pre-carbonate contamination) mantle melt, suggesting its ultrabasic nature. The reaction we envisage can be simplified as follows:



Equation (9) was proposed by Walter (ref.⁸²) and can explain the reaction observed in the rim of olivine phenocrysts in contact with monticellite. Another possible reaction could be represented by Bowen (ref.⁸³):



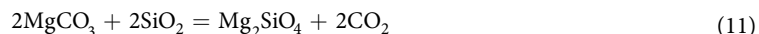
In their experiment (run 486) carried out in the $\text{CaO-MgO-SiO}_2\text{-CO}_2\text{-H}_2\text{O}$ system at 1038 °C and 2 kb, Otto and Wyllie (ref.⁸⁴; p. 357) report “clusters of subhedral crystals of monticellite with enclosed euhedral crystals of forsterite” associated with subhedral mostly angular calcite crystals and ~20 wt% SiO_2 in the experimental run. This paragenesis and the SiO_2 content closely resemble that of Polino rocks, and is, therefore, compatible with a process of monticellite crystallization at shallow depths rather than invoking equilibrium with forsterite at $P > 1$ GPa, as instead proposed by literature models. In the Otto and Wyllie (ref.⁸⁴) experiments, forsterite is always the major liquidus silicate mineral in low- SiO_2 mixtures. With increasing SiO_2 in the starting material, the liquidus minerals are olivine (melt $\text{SiO}_2 = 5\text{--}10$ wt%), olivine and monticellite (melt $\text{SiO}_2 = 15\text{--}20$ wt%), olivine and åkermanite or monticellite and åkermanite (melt $\text{SiO}_2 = 25\text{--}30$ wt%), åkermanite and diopside (melt $\text{SiO}_2 = 30\text{--}35$ wt%), diopside (melt $\text{SiO}_2 = 40$ wt%) and diopside and quartz (melt $\text{SiO}_2 = 45\text{--}50$ wt%; ref.⁸⁴).

Carbonatite Melt-Silicate Solid vs. Silicate Melt-Carbonatite Solid Constraints

The data presented above and those from previous experimental studies suggest a different origin for the CaO enrichment in Polino rocks, i.e. a shallow depth provenance rather than an upper mantle source. The strongly radiogenic $^{87}\text{Sr}/^{86}\text{Sr}$ (>0.710 ; ref.⁸⁵) and unradiogenic $^{143}\text{Nd}/^{144}\text{Nd}$ (<0.5122) of Polino and other ULUD carbonatites and calcite separates overlap the ULUD silicatic rocks. These values are far from the field of worldwide carbonatites ($^{87}\text{Sr}/^{86}\text{Sr} < 0.7065$; $^{143}\text{Nd}/^{144}\text{Nd} > 0.5123$; refs.^{47,85,86}), questioning the interpretation of Polino calcite as a mantle carbonatitic component. The $^{87}\text{Sr}/^{86}\text{Sr}$ of Polino calcite and whole-rock carbonatite is also higher than that of local Calcare Massiccio limestone (<0.7075 ; ref.⁸⁵), but this cannot be interpreted against a sedimentary origin of the Polino carbonate fraction. Indeed, in a hypothetical limestone-silicatic magma interaction, the Sr-rich ULUD magma (Sr up to 4000 ppm) strongly controls the variation of $^{87}\text{Sr}/^{86}\text{Sr}$ ratios in the contaminated melts. In this case, minimum amounts of magma can easily modify the original $^{87}\text{Sr}/^{86}\text{Sr}$ of the carbonate-rich

fraction of the contaminated magma, being the sedimentary fraction characterized by more than one order of magnitude less Sr.

Dolostones are present in minor amount in the carbonate succession pierced by the magma and could be potentially responsible for the high Fo content of Polino olivines. According to Iacono Marziano *et al.* (ref.⁷⁵), the magnesite end-member of dolomite reacts with $\text{SiO}_{2(\text{melt})}$ via:



Dolomite digestion often results in crystallization of MgO-rich olivine together with clinopyroxene, this latter being the only crystallizing phase in calcite-doped experiments⁷⁵. The very low NiO of MgO-rich and CaO-rich composition of Alban Hills volcano olivines is considered not a mantle origin feature, but rather the effect of digestion of dolostone⁴³, a conclusion confirmed by experimental data⁷⁵. North Baikal olivines formed after digestion of dolostones by mafic magma are characterized by similar low NiO (<0.25 wt%; average 0.12 wt%)⁸⁷. Polino olivines are characterized by high NiO (~0.06–0.95 wt%; average ~0.36 wt%, ~3300 ppm Ni), without any correlation between NiO and MgO ($R^2 = 0.06$; Table A2). These and the previous considerations lead us to hypothesize for the Fo-rich composition of Polino olivines a mantle origin, rather than a process of interaction with dolostones. This conclusion is corroborated by the very homogeneous composition of the largest olivines (Table A2), indicating the absence of variation from the core to the near rim sectors. Only the outermost olivine rims show MgO increase, associated to the CaO increase, as consequence of interaction with limestone. This process leads Ca entering the M2 olivine site, leaving Mg free entering the M1 site, not preferred by Fe (ref.⁸¹).

As concerns the stable isotopes, the high $\delta^{18}\text{O}$ (~+24‰; ref.⁹) of the calcite component cannot be used to infer a mantle origin for the carbonate component in Polino rocks. The low $\delta^{13}\text{C}$ of Polino calcite (from –8.0 to –13.1‰; ref.⁹) is within the $\delta^{13}\text{C}$ values of Alban Hills skarns (from +3 to –12‰; ref.⁵³), Alban Hills lavas (from +5 to –19‰; ref.⁵³) and the Morron de Villamayor sedimentary calcite entrapped in olivine (from –11 to –12‰; ref.⁸⁸). Natural Alban Hills clinopyroxenes show good correlation between $\delta^{18}\text{O}$ and ^{IV}Al with increasing CaO content⁵¹. Experimental modelling of AFC process adding carbonate component to a silicatic magma shows the same compositional evolution of clinopyroxenes, with the highest kushiroitic component recorded in the experimental runs with higher amounts of added carbonates (e.g., refs^{78,81}). If these stable isotope data cannot be used to definitively assume a sedimentary origin for the carbonate fraction of Polino rocks, they are at least compatible with such an origin.

As monticellite is known to be unstable at pressures >1 GPa (ref.⁸⁹), the hypothesis of monticellite formation after the reaction of mantle minerals with Ca-carbonatitic melt below the Moho appears improbable, and implies its secondary origin. In other words, our view overturns the classically accepted interpretation^{8,9} of monticellite as the reaction product between a carbonatitic magma infiltrating mantle matrix, then scraped off by the same carbonatitic liquid. We believe, instead, that the interaction occurred between an ultrabasic melt and the sedimentary carbonate wall rocks *en route* to the surface.

The presence of phlogopite, forsterite, monticellite, perovskite and periclase (the first four phases identified in Polino rocks) is not a proof to discriminate a carbonate-rich rock as carbonatite, being these minerals present in impure marbles too^{38,50,90}. Also, the presence of schorlomite garnet in Polino rocks⁵⁹ cannot be considered a proof for a carbonatitic origin of the melt, being this mineral a typical phase of impure limestones, marbles and skarns (e.g., refs^{90–92}). According to several authors (refs^{29,38,43,51}), no decisive chemical and mineralogical criterions by which to establish a magmatic heritage for a carbonate-rich rock do exist, considering that also magmatic carbonatite can undergo drastic compositional changes after emplacement.

Recognition of primary carbonate components must be based on the association with igneous rocks, the presence of fenites, euhedral crystals of carbonates evidencing flow structures, the mineral paragenesis, the incompatible element composition and the radiogenic and stable isotopic ratios of the carbonate-rich rocks²⁷. None of these evidences can be found at Polino: (1) no silicate pair rocks are associated to the “carbonatitic” diatreme; (2) no fenitization is recorded on country rocks; (3) no euhedral tabular or rhombic crystals of calcite have been found; (4) no flow structures of carbonate materials have been identified; (5) monticellite stability field is constrained to shallow P (<1 GPa), rendering impossible the formation of this mineral at mantle depths; (6) monticellite appearance is contemporaneous with cessation of olivine growth and crystallization; (7) the incompatible and compatible element budget of Polino “carbonatite” is different from average worldwide Ca-carbonatites^{62,63,93} and Fort Portal monticellite carbonatite⁶⁸; (8) Radiogenic isotopic ratios indicate interaction with crustal lithologies, while stable isotopic ratios are compatible with C-O exchanges at relatively low-T conditions.

Conclusions

At Polino, plagioclase- and pyroxene-free rocks characterized by abundant Fo-rich olivine, monticellite, calcite, phlogopite, perovskite, Fe-Ti oxides were emplaced ~246 ka. Their whole-rock composition is intermediate between Ca-carbonatitic magma and alkali-poor ultrabasic melt.

On the basis of petrographic evidences (euhedral shape of olivine and phlogopite crystals, presence of monticellite as groundmass phase and as rim around olivine), mineral chemical compositions (homogeneous Fo-rich and modestly CaO-enriched olivine, high Mg# phlogopite), experimental constraints (stability field of monticellite limited to <1 GPa), isotopic data (strongly radiogenic $^{87}\text{Sr}/^{86}\text{Sr}$, high $\delta^{18}\text{O}$, low $\delta^{13}\text{C}$) and geological-volcanological considerations (small volume of the magma, limited diameter of the diatreme, over-thickened sedimentary carbonate country-rocks pierced by the upwelling magma), we propose that Polino rocks cannot be considered as true carbonatites (monticellite alvikite), as instead reported in literature. The origin of these rocks is, rather, more compatible with interaction between an ultrabasic melt likely generated by a magnesite-bearing residual peridotite and local sedimentary carbonate rocks.

The classification of Polino volcanic rocks remains debated. It cannot be defined as alvikite (too low primary carbonate content), basalt (no plagioclase), kamafugite (no kalsilite), melilitite (no melilite), foidite (no foids), dunite (not a plutonic/metamorphic rock) or picrite ($\text{SiO}_2 < 30 \text{ wt}\%$). It represents an alkali-poor strongly ultrabasic melt whose original composition was modified by the digestion of sedimentary carbonates, and can consequently be classified as pseudocarbonatite³⁰. The final message is that the carbonate component of Polino rocks is not mantle-derived, but rather is the result of partial digestion of shallow limestones.

On the other hand, the original Polino magma (i.e., the composition before interaction with sedimentary carbonates) is considered to derive from a carbonated (magnesite-bearing) peridotitic source, to explain the ultrabasic compositions (absence of feldspars and clinopyroxene) and the Mg-rich olivine and phlogopite phenocrysts. In other words, the petrogenetic model proposed is a fuzzy one, based not on a carbonatite-yes or carbonatite-no choice. At the same time, the Polino rocks are not true carbonatites, but they derive from a carbonated mantle source.

References

1. Stoppa, F. & Lavecchia, G. Late Pleistocene ultra-alkaline magmatic activity in the Umbria-Latium region Italy: an overview. *J. Volcanol. Geotherm. Res.* **52**, 277–293 (1992).
2. Carminati, E., Lustrino, M. & Doglioni, C. Geodynamic evolution of the central and western Mediterranean: Tectonics vs. igneous petrology constraints. *Tectonophysics* **579**, 173–192 (2012).
3. Paone, A. A Review of Carbonatite Occurrences in Italy and Evaluation of Origins. *Open J. Geol.* **3**, 66–82 (2013).
4. Peccerillo, A. The Intra-Appennine Province. In: *Cenozoic Volcanism in the Tyrrhenian Sea Region*. (Springer, Cham) pp. 61–79 (2017).
5. Le Maitre, R. W. (Ed.) *Igneous Rocks. A Classification and Glossary of Terms. Recommendations of the International Union of Geological Sciences Subcommittee on the Systematics of Igneous Rocks, 2nd ed.* (Cambridge Univ. Press, New York, 2002).
6. Laurenzi, M., Stoppa, F. & Villa, I. Eventi ignei monogenici e depositi piroclastici nel distretto ultra-alciano umbro-laziale ULUD: revisione, aggiornamento e comparazione dei dati cronologici. *Plinius* **12**, 61–65 (1994).
7. Stoppa, F. & Lupini, L. Caratteristiche identificative di una roccia carbonatitica del Pleistocene superiore affiorante presso Polino (Tr-Umbria). *Stud. Geol. Camerti Vol. Spec.* 383–398 (1991).
8. Stoppa, F. & Lupini, L. Mineralogy and petrology of the Polino Monticellite Calcicarbonatite (Central Italy). *Mineral. Petrol.* **49**, 213–231 (1993).
9. Rosatelli, G., Wall, F., Stoppa, F. & Brilli, M. Geochemical distinctions between igneous carbonate, calcite cements, and limestone xenoliths (Polino carbonatite, Italy): Spatially resolved LAICPMS analyses. *Contrib. Mineral. Petrol.* **160**, 645–661 (2010).
10. Martin, L. H. J. *et al.* Element partitioning between immiscible carbonatite-kamafugite melts with application to the Italian ultrapotassic suite. *Chem. Geol.* **320**, 96–112 (2012).
11. Lavecchia, G., Creati, N. & Boncio, P. The intramontane ultra-alkaline Province (IUP) of Italy: a brief review with considerations on the thickness of the underlying lithosphere. *Boll. Soc. Geol. Ital. Vol. Spec.* **1**, 87–98 (2002).
12. Locardi, E. & Nicolich, R. Crust-mantle structures and Neogene-Quaternary magmatism in Italy. *Boll. Geofis. Teor. Appl.* **46**, 169–180 (2005).
13. Bell, K., Lavecchia, G. & Rosatelli, G. Cenozoic Italian magmatism - Isotope constraints for possible plume-related activity. *J. South Am. Earth Sci.* **41**, 22–40 (2013).
14. Kogarko, L. N. & Veselovskiy, R. V. Geodynamic origin of carbonatites from the absolute paleotectonic reconstructions. *J. Geodyn.* **125**, 13–21 (2019).
15. Cundari, A. Role of subduction in the genesis of potassic basaltic rocks: a discussion paper on the unfashionable side of the role. *Miner. Petrogr. Acta* **37**, 121–130 (1994).
16. Peccerillo, A. Relationships between ultrapotassic and carbonate-rich volcanic rocks in central Italy: petrogenetic and geodynamic implications. *Lithos* **43**, 267–279 (1998).
17. Melluso, L., Conticelli, S., D'Antonio, M., Mirco, N. P. & Saccani, E. Petrology and mineralogy of wollastonite- and melilite-bearing paralavas from the Central Apennines. *Am. Mineral.* **88**, 1287–1299 (2003).
18. Peccerillo, A. & Lustrino, M. Compositional variations of the Plio-Quaternary magmatism in the circum-Tyrrhenian area: deep versus shallow-mantle processes. *Geol. Soc. Am. Spec. Paper* **388**, 421–434 (2005).
19. Avanzinelli, R., Lustrino, M., Mattei, M., Melluso, L. & Conticelli, S. Potassic and ultrapotassic magmatism in the circum-Tyrrhenian region: Significance of carbonated pelitic vs. pelitic sediment recycling at destructive plate margins. *Lithos* **113**, 213–227 (2009).
20. Conticelli, S. *et al.* Leucite-bearing (kamafugitic/leucitic) and -free (lamproitic) ultrapotassic rocks and associated shoshonites from Italy: Constraints on petrogenesis and geodynamics. *J. Virtual Explor.* **36**, 20, <https://doi.org/10.3809/jvirtex.2010.00251> (2010).
21. Lustrino, M., Duggen, S. & Rosenberg, C. L. The Central-Western Mediterranean: anomalous igneous activity in an anomalous collisional tectonic setting. *Earth Sci. Rev.* **104**, 1–40 (2011).
22. D'Orazio, M., Innocenti, F., Tonarini, S. & Doglioni, C. Carbonatites in a subduction system: The Pleistocene alvikites from Mt. Vulture (southern Italy). *Lithos* **98**, 313–334 (2007).
23. Bailey, D. K. Lithosphere control of continental rift magmatism. *J. Geol. Soc. London* **133**, 103–106 (1977).
24. Foley, S. F. & Fischer, T. B. An essential role for continental rifts and lithosphere in the deep carbon cycle. *Nature Geosci.* **10**, 897–902 (2017).
25. Woolley, A. R. & Kjarsgaard, B. A. Paragenetic types of carbonatite as indicated by the diversity and relative abundances of associated silicate rocks: evidence from a global database. *Can. Mineral.* **46**, 741–752 (2008).
26. Canil, D. Experimental study bearing on the absence of carbonate in mantle-derived xenoliths. *Geology* **18**, 1011–1013 (1990).
27. Ionov, D. & Harmer, R. E. Trace element distribution in calcite-dolomite carbonatites from Spitskop: inferences for differentiation of carbonatite magmas and the origin of carbonates in mantle xenoliths. *Earth Planet. Sci. Lett.* **198**, 495–510 (2002).
28. Bianchini, G. & Natali, C. Carbon elemental and isotopic composition in mantle xenoliths from Spain: Insights on sources and petrogenetic processes. *Lithos* **272–273**, 84–91 (2017).
29. Barker, D. S. Field relations of carbonatites In: Bell, K. (Ed.), *Carbonatites: Genesis and Evolution*, Unwin Hyman, London, 38–69 (1989).
30. Woolley, A. R. & Church, A. A. Extrusive carbonatites: a brief review. *Lithos* **85**, 1–14 (2005).
31. Mitchell, R. H. Carbonatites and carbonatites and carbonatites. *Can. Mineral.* **43**, 2049–2068 (2005).
32. Harmer, R. E. & Gittins, J. The Case for Primary, Mantle-derived Carbonatite Magma. *J. Petrol.* **39**, 1895–1903 (1998).
33. Franz, G. & Wyllie, P. Experimental Studies in the System $\text{CaO-MgO-SiO}_2\text{-CO}_2\text{-H}_2\text{O}$. In *Ultramafic and related rocks* (ed. Wyllie, P. J.) 323–326 (1967).
34. Freestone, I. C. & Hamilton, D. L. The role of liquid immiscibility in the genesis of carbonatites - An experimental study. *Contrib. Mineral. Petrol.* **73**, 105–117 (1980).

35. Kjarsgaard, B. A. & Hamilton, D. L. The genesis of carbonatites by liquid immiscibility. In *Carbonatites: Genesis and Evolution* (ed. Bell, K.), Unwin Hyman, London, 388–404 (1989).
36. Brooker, R. A. & Kjarsgaard, B. A. Silicate-carbonate liquid immiscibility and phase relations in the system $\text{SiO}_2\text{-Na}_2\text{O-Al}_2\text{O}_3\text{-CaO-CO}_2$ at 0.1–2.5 GPa with applications to carbonatite genesis. *J. Petrol.* **52**, 1281–1305 (2011).
37. Ngwenya, B. Hydrothermal rare earth mineralisation in carbonatites of the Tundulu complex, Malawi: Processes at the fluid/rock interface. *Geochim. Cosmochim. Acta* **58**, 2061–2072 (1994).
38. Lentz, D. R. Carbonatite genesis: a reexamination of the role of intrusion-related pneumatolytic skarn processes in limestone melting. *Geology* **27**, 335–338 (2009).
39. Gudfinnsson, G. H. & Presnall, D. C. Continuous Gradations among Primary Carbonatitic, Kimberlitic, Melilititic, Basaltic, Picritic, and Komatiitic Melts in Equilibrium with Garnet Lherzolite at 3–8 GPa. *J. Petrol.* **46**, 1645–1659 (2005).
40. Hammouda, T. & Keshav, S. Melting in the mantle in the presence of carbon: review of experiments and discussion on the origin of carbonatites. *Chem. Geol.* **418**, 171–188 (2015).
41. Lustrino, M. What ‘anorogenic’ igneous rocks can tell us about the chemical composition of the upper mantle: case studies from the circum-Mediterranean area. *Geol. Mag.* **148**, 304–316 (2011).
42. Mariano, A. N. & Roeder, P. L. Kerimasi: A Neglected Carbonatite Volcano. *J. Geol.* **91**, 449–455 (1983).
43. Federico, M., Peccerillo, A., Barbieri, M. & Wu, T. W. Mineralogical and geochemical study of granular xenoliths from the Alban Hills volcano, Central Italy: bearing on evolutionary processes in potassic magma chambers. *Contrib. Mineral. Petrol.* **115**, 384–401 (1994).
44. Lavecchia, G. & Stoppa, F. Tyrrhenian zone as intra-continental magmatism. *Earth Planet. Sci. Lett.* **99**, 336–350 (1990).
45. Stoppa, F. & Woolley, A. R. The Italian carbonatites: Field occurrence, petrology and regional significance. *Mineral. Petrol.* **59**, 43–67 (1997).
46. Barbieri, M., D’Orefice, M., Graciotti, R. & Stoppa, F. Il vulcanismo monogenico medio-pleistocenico della conca di Carsoli (L’Aquila). *Geol. Rom.* **36**, 13–31 (2001).
47. Bianchini, G., Beccaluva, L. & Siena, F. A reappraisal of ultra-alkaline Intra-Apennine volcanism in central southern Italy: evidence for subduction-modified mantle sources in Special Issue: A showcase of the Italian research in petrology: magmatism in Italy. *Per. Mineral.* **73**, 177–185 (2004).
48. Rosatelli, G., Stoppa, F. & Jones, A. P. Intrusive calcite-carbonatite occurrence from Mt. Vulture volcano, southern Italy. *Mineral. Mag.* **64**, 615–624 (2000).
49. Downes, H. *et al.* Geochemistry and Sr-Nd isotopic compositions of mantle xenoliths from the Monte Vulture carbonatite-melilitite volcano, central southern Italy. *Contrib. Mineral. Petrol.* **144**, 78–92 (2002).
50. Freda, C., Gaeta, M., Palladino, D. M. & Trigila, R. The Villa Senni Eruption (Alban Hills, central Italy): The role of H_2O and CO_2 on the magma chamber evolution and on the eruptive scenario. *J. Volcanol. Geotherm. Res.* **78**, 103–120 (1997).
51. Federico, M. & Peccerillo, A. Mineral chemistry and petrogenesis of granular ejecta from the Alban Hills volcano (Central Italy). *Mineral. Petrol.* **74**, 223–252 (2002).
52. Dallai, L., Freda, C. & Gaeta, M. Oxygen isotope geochemistry of pyroclastic clinopyroxene monitors carbonate contributions to Roman-type ultrapotassic magmas. *Contrib. Mineral. Petrol.* **148**, 247–263 (2004).
53. Di Rocco, T., Freda, C., Gaeta, M., Mollo, S. & Dallai, L. Magma chambers emplaced in carbonate substrate: Petrogenesis of skarn and cumulate rocks and implications for CO_2 degassing in volcanic areas. *J. Petrol.* **53**, 2307–2332 (2012).
54. Gozzi, F. *et al.* Primary magmatic calcite reveals origin from crustal carbonate. *Lithos* **190–191**, 191–203 (2014).
55. Shaw, C. S. J. Evidence for the presence of carbonate melt during the formation of cumulates in the Colli Albani Volcanic District, Italy. *Lithos* **310–311**, 105–119 (2018).
56. Stoppa, F., Lloyd, F. & Rosatelli, G. CO_2 as the virtual propellant of carbonatite-kamafugite conjugate pairs and the eruption of diatremic tuffsite. *Per. Mineral.* **72**, 205–222 (2003).
57. Peccerillo, A. Carbonate-rich pyroclastic rocks from central Apennines: carbonatites or carbonated rocks? A commentary. *Per. Mineral.* **73**, 165–175 (2004).
58. Brozzetti, F., Lavecchia, G. & Stoppa, F. Presentazione di alcuni aspetti tettonico-strutturali e petrologici del distretto ultra-alkalino umbro-laziale. *Studi Geol. Camerti CROP* **11**, 327–337 (1991).
59. Lupini, L., Williams, C. T. & Woolley, A. R. Zr-rich garnet and Zr- and Th-rich perovskite from the Polino carbonatite, Italy. *Mineral. Mag.* **56**, 581–586 (1992).
60. Brandano, M., Corda, L., Tomassetti, L. & Tagliavento, M. Frequency analysis across the drowning of a Lower Jurassic carbonate platform: The Calcare Massiccio Formation (Apennines, Italy). *Mar. Pet. Geol.* **78**, 606–620 (2016).
61. Lustrino, M. *et al.* Petrogenesis of Early Cretaceous silicic volcanism in SE Uruguay: The role of mantle and crustal sources. *Geochem. J.* **44**, 1–22 (2010).
62. Dalton, J. A. & Wood, B. J. The compositions of primary carbonate melts and their evolution through wallrock reaction in the mantle. *Earth Planet. Sci. Lett.* **119**, 511–525 (1993).
63. Lee, W. J. & Wyllie, P. J. Experimental data bearing on liquid immiscibility, crystal fractionation, and the origin of calciocarbonatites and natrocarbonatites. *Int. Geol. Rev.* **36**, 797–819 (1994).
64. Bailey, D. K. & Collier, J. D. Carbonatite-melilitite association in the Italian collision zone and the Ugandan rifted craton: significant common factors. *Mineral. Mag.* **64**, 675–682 (2000).
65. Woolley, R. A. & Kempe, D. R. C. Carbonatites: Nomenclature, average chemical composition. In *Carbonatites: Genesis and Evolution* (ed. Bell, K.), Unwin Hyman, London, 1–14 (1989).
66. Chakhmouradian, A. R. High-field-strength elements in carbonatitic rocks: Geochemistry, crystal chemistry and significance for constraining the sources of carbonatites. *Chem. Geol.* **235**, 138–160 (2006).
67. Bell, K. & Kjarsgaard, B. Discussion of Peccerillo (2004) “Carbonate-rich pyroclastic rocks from central Apennines: carbonatites or carbonated rocks?” *Per. Mineral.* **75**, 85–92 (2006).
68. Eby, G. N., Lloyd, F. E. & Woolley, A. R. Geochemistry and petrogenesis of the Fort Portal, Uganda, extrusive carbonatite. *Lithos* **113**, 785–800 (2009).
69. Barker, D. S. & Nixon, P. H. High-Ca, low-alkali carbonatite volcanism at Fort Portal, Uganda. *Contrib. Mineral. Petrol.* **103**, 166–177 (1989).
70. Nixon, P. H. (Ed.) *Mantle Xenoliths*. John Wiley Ed., 844 pp (1987).
71. Ammannati, E., Jacob, D. E., Avanzinelli, R., Foley, S. F. & Conticelli, S. Low Ni olivine in silica-undersaturated ultrapotassic igneous rocks as evidence for carbonate metasomatism in the mantle. *Earth Planet. Sci. Lett.* **444**, 64–74 (2016).
72. Sobolev, A. V., Hofmann, A. W., Sobolev, S. V. & Nikogosian, I. K. An olivine-free mantle source of Hawaiian shield basalts. *Nature* **434**, 590–597 (2005).
73. Lustrino, M. How the delamination and detachment of lower crust can influence basaltic magmatism. *Earth-Sci. Rev.* **72**, 21–38 (2005).
74. Roeder, P. L. & Emslie, R. F. Olivine-liquid equilibrium. *Contrib. Mineral. Petrol.* **29**, 275–289 (1970).
75. Iacono Marziano, G., Gaillard, F. & Pichavant, M. Limestone assimilation and the origin of CO_2 emission at the Alban Hills (Central Italy): Constraints from experimental petrology. *J. Volcanol. Geotherm. Res.* **166**, 91–105 (2007).
76. Barchi, M., Minelli, G., Magnani, B. & Mazzotti, A. Line CROP 03: Northern Apennines. *Mem. Descr. Carta Geol. It.* **62**, 127–136 (2003).

77. Iacono Marziano, G., Gaillard, F. & Pichavant, M. Limestone assimilation by basaltic magmas: an experimental re-assessment and application to Italian volcanoes. *Contrib. Mineral. Petrol.* **155**, 719–738 (2008).
78. Mollo, S. *et al.* Carbonate assimilation in magmas: A reappraisal based on experimental petrology. *Lithos* **114**, 503–514 (2010).
79. Mollo, S. & Vona, A. The geochemical evolution of clinopyroxene in the Roman Province: A window on decarbonation from wall-rocks to magma. *Lithos* **192**, 1–7 (2014).
80. Carter, L. B. & Dasgupta, R. Hydrous basalt-limestone interaction at crustal conditions: Implications for generation of ultracalcic melts and outflux of CO₂ at volcanic arcs. *Earth Planet. Sci. Lett.* **427**, 202–214 (2015).
81. Di Stefano, F. *et al.* Olivine compositional changes in primitive magmatic skarn environments: A reassessment of divalent cation partitioning models to quantify the effect of carbonate assimilation. *Lithos* **316–317**, 104–121 (2018).
82. Walter, L. S. Experimental studies on Bowen's decarbonation series: II: P-T univariant equilibria of the reaction: forsterite + calcite = monticellite + periclase + CO₂. *Am. Mineral.* **261**, 773–779 (1963).
83. Bowen, N. L. Progressive Metamorphism of Siliceous Limestone and Dolomite. *J. Geol.* **48**, 225–274 (1940).
84. Otto, J. W. & Wyllie, P. J. Relationships between silicate melts and carbonate-precipitating melts in CaO-MgO-SiO₂-CO₂-H₂O at 2 kbar. *J. Petrol.* **48**, 343–365 (1993).
85. Castorina, F., Stoppa, F., Cundari, A. & Barbieri, M. An enriched mantle source for Italy's melilitite-carbonatite association as inferred by its Nd-Sr isotope signature. *Mineral. Mag.* **64**, 625–639 (2000).
86. Gomes, C. B., Comin-Chiaromonte, P., Azzone, R. G., Ruberti, E. & Rojas, G. E. E. Cretaceous carbonatites of the southeastern Brazilian platform: a review. *Braz. J. Geol.* **48**, 317–345 (2018).
87. Wenzel, T., Baumgartner, L. P., Brugmann, G. E., Konnikov, E. G. & Kislov, E. V. Partial melting and assimilation of dolomitic xenoliths by mafic magma: the Ioko-Dovyren intrusion (North Baikal Region, Russia). *J. Petrol.* **43**, 2049–2074 (2002).
88. Lustrino, M. *et al.* Ca-rich carbonates associated with ultrabasic-ultramafic melts: Carbonatite or limestone xenoliths? A case study from the late Miocene Morron de Villamayor volcano (Calatrava Volcanic Field, central Spain). *Geochim. Cosmochim. Acta* **185**, 477–497 (2016).
89. Liu, L. G. The high-pressure phase transformations of monticellite and implications for upper mantle mineralogy. *Phys. Earth Planet. Int.* **20**, 25–29 (1979).
90. Huckenholz, H. G. Synthesis and stability of Ti-andradite. *Am. J. Sci.* **267**, 209–232 (1969).
91. Gittins, J. Some observations on the present status of carbonatite studies. *Proc. First Int. Symp. Carbonatites* 107–115 (1978).
92. Chakmouradian, A. R. & McCammon, C. A. Schorlomite: A discussion of the crystal chemistry, formula, and inter-species boundaries. *Phys. Chem. Miner.* **32**, 277–289 (2005).
93. Kapustin, Y. L. Geochemistry of strontium and barium in carbonatites. *Geokhimiya* **3**, 369–380 (1982).

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Author Contributions

M.L. conceived the research project, proposed the petrogenetic model, wrote the main text and drafted the figures. M.L. and V.S. reviewed and edited the manuscript. M.L., N.L. and V.S. collected the rocks, performed the petrographic description, and conducted EMP and SEM analyses.

Additional Information

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