

1 TITLE:

2 **EFFECTS OF COSEISMIC GROUND VERTICAL MOTION ON MASONRY**  
3 **CONSTRUCTIONS DAMAGE DURING THE 2016 AMATRICE-NORCIA (CENTRAL**  
4 **ITALY) EARTHQUAKES**

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22 ABSTRACT

23

24 Horizontal acceleration and velocity are considered the most important parameters in determining  
25 damage potential to buildings during the paroxysmal energy dissipation of an earthquake. However,  
26 taking as example the two mainshocks of the 2016 Central Italy seismic sequence and comparing  
27 Interferometric Synthetic-Aperture Radar (InSAR) and macroseismic data, it is shown that maximum  
28 damage was concentrated where the ground subsided coseismically. A number of empirical  
29 relationships are determined between InSAR vertical displacement on the one hand, and  
30 macroseismic intensity or ground motion intensity measures on the other. Finally, resorting to a finite-  
31 discrete element model, the effect of the vertical component of ground motion is investigated on a  
32 test masonry structure under recorded accelerograms.

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35 KEYWORDS:

36 InSAR; macroseismic intensity; intensity measures; finite-discrete element.

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39 1 INTRODUCTION

40

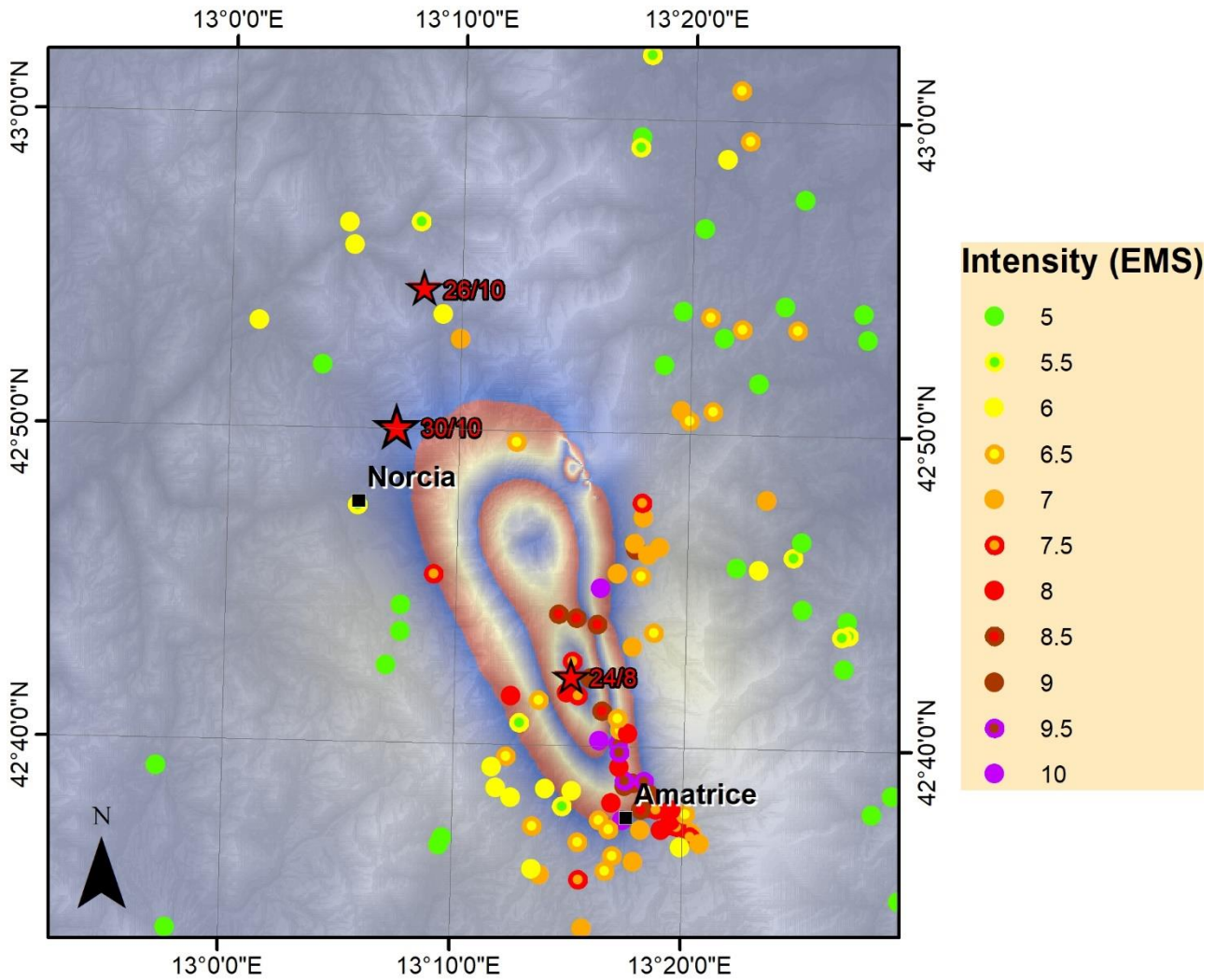
41 In 2016, Central Italy was struck by a seismic sequence consisting of two mainshocks,  $M_w$  6.0 on  
42 August 24th and  $M_w$  6.5 on October 30th (Chiaraluce et al. 2017). The seismicity was generated by  
43 the tectonic extensional stress regime affecting the Apennines ridge (Bigi et al., 1989, Petricca et al.  
44 2015, Doglioni et al. 2015), where a rate of around N50°E dilation of about 4-5 mm/yr is recorded  
45 through permanent GPS networks (Devoti et al. 2017) across the belt from the Tyrrhenian coast to  
46 the highest peaks, which are on average slightly east of the water divide. A large number of  
47 multidisciplinary data were collected during the sequence, namely seismological, geological,  
48 geodetic (GPS and Interferometric Synthetic-Aperture Radar (InSAR)), strong motion and  
49 macroseismic surveys (Cheloni et al. 2017, Wilkinson et al. 2017, Azzaro et al. 2016, Tertulliani and  
50 Azzaro 2016). This research was motivated by the overlap between the most damaged areas and the  
51 coseismically subsided zone. The concentration of damage can be explained locally by site-  
52 amplification effects and seismic wave directivity (Azzaro et al. 2016, Calderoni et al. 2017).  
53 However, a further effect could be provided by the combination of the horizontal components of  
54 ground motion and the simultaneous vertical acceleration (Ganz and Doglioni 2014, Mariani and Pugi  
55 2018). In this study, InSAR vertical displacement is compared with macroseismic data and ground  
56 motion intensity in order to explain the greater destruction of masonry buildings in the coseismically  
57 subsided area (Figures 1-2). In Figure 2, macroseismic data also cumulate the effects of the previous  
58  $M_w$  6.0 August 24 earthquake. It is evident from interferometric data that the area struck by the  
59 earthquake is divided into two coseismically deformed areas, i.e., subsided and uplifted. The subsided  
60 area is much more pronounced than them part that was uplifted, and macroseismic data point to more  
61 severe damage within the subsided area. In fact, the most severe damage is concentrated in the area  
62 that underwent subsidence, whereas settlements closer to the epicentre, but coseismically uplifted,  
63 experienced less destruction.

64 This study investigates the influence of combined horizontal and vertical components (Figure 3) on  
65 the dynamic nonlinear response of unreinforced masonry structures. Downward (upward)  
66 acceleration entails instantaneous decrement (increment) of friction force and corresponding potential

67 increment (decrement) of sliding between the bricks or stones of masonry buildings, depending on  
68 the level of horizontal action at the same instant.

69 This article focuses on this observation, investigating numerically the dynamic effect of the variation  
70 of friction forces between units and mortar induced by vertical acceleration.

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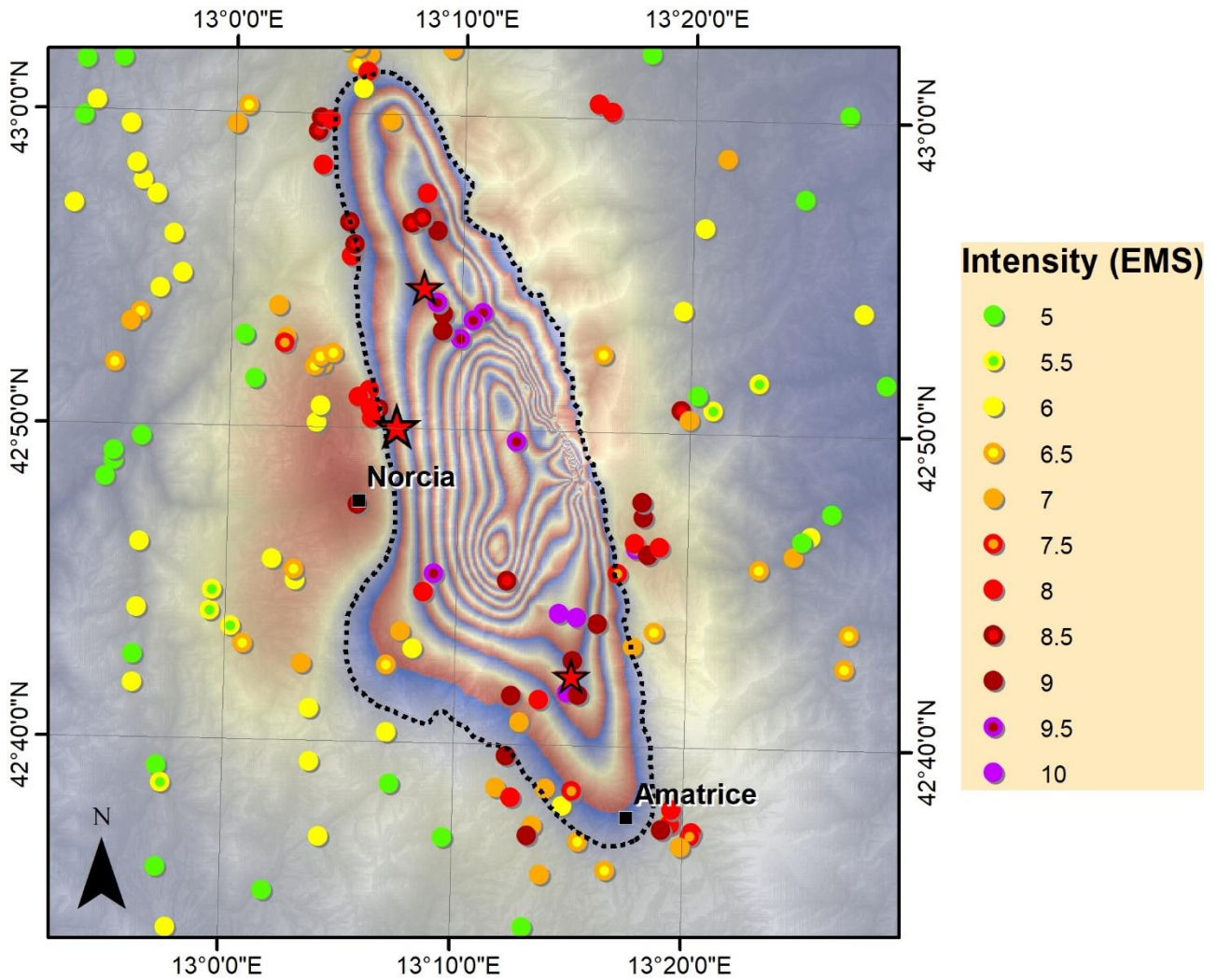


72

73 Figure 1. InSAR coseismic subsidence (max ~200 mm) and macroseismic data associated with the  
74  $M_w$  6.0 August 24th 2016 mainshock. The vertical displacement field has been rewrapped so that  
75 each fringe corresponds to 100 mm.

76

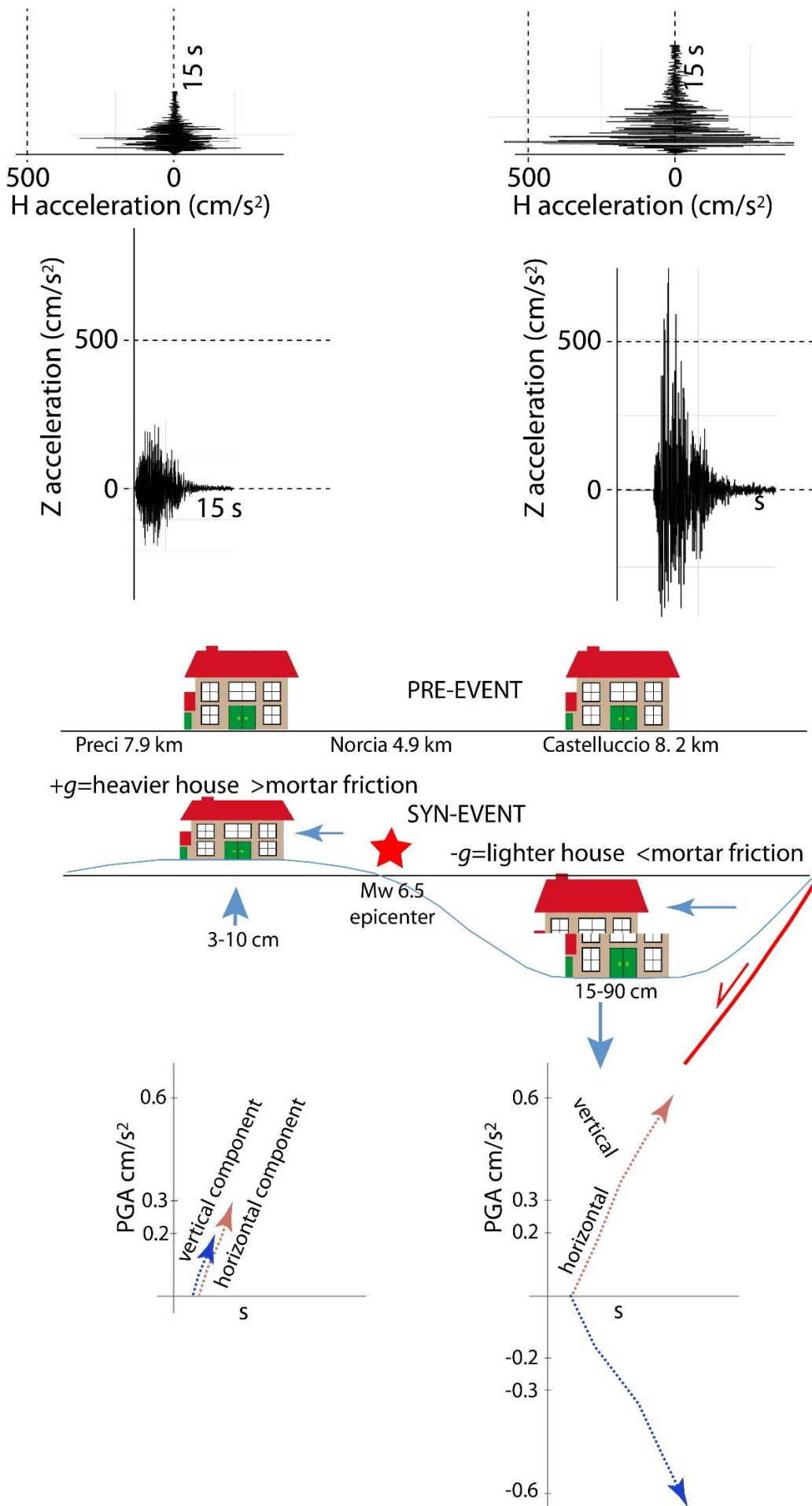
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78

79 Figure 2. InSAR coseismic subsidence (max ~1000 mm east of Norcia) and macroseismic data  
80 associated with the  $M_w$  6.5 October 30th 2016 mainshock. The vertical displacement field has been  
81 rewrapped so that each fringe corresponds to 100 mm.

82



84 Figure 3. Model investigated in this study. Higher panels: Ground motion components occurring at  
85 the same time. Lower panels: A short fraction of an accelerogram, with acceleration versus time (s)  
86 showing two cases in which the coseismic motion is upward (left) or downward (right) during the  
87 increase of horizontal components.

88

89

## 90 2 InSAR DATA

91

92 Displacement data used in this analysis are derived from InSAR measurements. InSAR processing  
93 is, nowadays, one of the most commonly used techniques to depict the spatial distribution of the  
94 permanent displacement occurred after a seismic event (see Weston et al. 2011 and references  
95 therein). This technique is based on radar phase comparison of two SAR images acquired with the  
96 same satellite, at the same orbit and at the same position, but at two distinct times, before and after  
97 the event (Franceschetti and Lanari 1999).

98 Though the description of the InSAR rationale is beyond the goal of this work, it is worth noting that  
99 the displacement measurements obtained from the radar phase difference are in the Line-of-Sight  
100 (LoS) direction, i.e. a diagonal line connecting the satellite to the pixel on the ground, nearly East-  
101 West and tilted  $\sim 40^\circ$  from the vertical (this angle has a rather wide range of values, according to the  
102 satellite). This peculiar SAR geometry makes the rigorous derivation of horizontal and vertical  
103 displacement components complex; in order to bypass the unavoidable approximations affecting  
104 every approach based only on the combination of different LoSs, the vertical displacement is derived  
105 from the Central Italy source models shown in Cheloni et al. (2017).

106 At the basis of this approach is analytical source modelling, whose rationale is described in Atzori et  
107 al. (2009): an exhaustive dataset of InSAR maps is handled by means of linear and non-linear  
108 inversion optimizations to derive the seismic sources responsible for the permanent displacement of  
109 both events. These sources are then used as input for forward modelling to get the estimated  
110 displacement along the East, North and vertical directions, the latter being the one used in this work.  
111 This model-based approach to get a complete 3D displacement has the further advantage of filtering  
112 out most of the atmospheric artifacts that generally affect InSAR maps and cannot be *a priori*  
113 discriminated from the real displacement signal. In fact, atmospheric artifacts are almost impossible  
114 to fit within the modelling optimization and are therefore absent in the forward calculation.

115 In addition to InSAR data used to get the source models for the August 24 and October 30  
116 earthquakes, and reported in Table 1, GPS data from the Ca.Geo.Net. network (Galvani et al. 2012)  
117 and the Italian Istituto Geografico Militare network were also considered.

118

119 Table 1. SAR image acquisitions.

SAR images		Satellite/ Constellation	Space Agency	Orbit	Radar band
pre-event	post-event				
<b>August, 24 event</b>					
August, 20	August, 28	COSMO-SkyMed	Italian (ASI)	ascending	X-band
August, 20	August, 26	COSMO-SkyMed	Italian (ASI)	descending	X-band
August, 21	August, 27	Sentinel-1	European (ESA)	ascending	C-band
August, 21	August, 27	Sentinel-1	European (ESA)	descending	C-band
September, 15*	August, 24	ALOS-2	Japanese (JAXA)	ascending	L-band
May, 25	August, 31	ALOS-2	Japanese (JAXA)	descending	L-band
<b>October, 30 event**</b>					
September, 9*	November, 2	ALOS-2	Japanese (JAXA)	ascending	L-band
May, 25	November, 9	ALOS-2	Japanese (JAXA)	descending	L-band

120 \* 2015

121 \*\* Both pairs include the effects of the August 24th event.

122

123

### 124 3 RELATIONSHIPS BETWEEN InSAR VERTICAL DISPLACEMENT AND 125 MACROSEISMIC INTENSITY

126

127 A first group of empirical relationships between InSAR vertical displacement and macroseismic  
128 intensity was investigated for the 2016 August 24 mainshock. Concerning the following events,  
129 although InSAR displacement was somewhere higher, like the one on the October 30, 2016, the  
130 determination of these relationships is more problematic because of cumulated intensity.

131 The Mercalli-Cancani-Sieberg (MCS) intensities,  $I_{MCS}$  (Galli et al. 2016) and the European  
132 Macroseismic Scale (EMS) intensities,  $I_{EMS}$  (Azzaro et al. 2016) are reported in Table 2 for the sites  
133 with  $I_{MCS}, I_{EMS} \geq VI$ , with the corresponding modelled InSAR vertical displacements, denoted by  $w$ .

134

135 Table 2.  $I_{MCS}, I_{EMS}$  vs. InSAR vertical displacement  $w$ , August 24th, 2016 mainshock ( $I_{MCS}, I_{EMS} \geq$   
136 VI).

Site	Municipality and Province abbreviation	$I_{MCS}$	$I_{EMS}$	$w$ (m)
------	---	-----------	-----------	---------

Amatrice	Amatrice RI	X-XI	X	-0.0464
Petrana	Amatrice RI	X-XI	IX-X	-0.0819
Pescara del Tronto	Arquata del Tronto AP	X-XI	X	-0.0764
Illica	Accumoli RI	X	IX-X	-0.1984
Casale	Amatrice RI	X	IX-X	-0.0948
Saletta	Amatrice RI	X	X	-0.1492
Rio	Amatrice RI	IX-X		-0.0870
San Lorenzo e Flaviano	Amatrice RI	IX-X	IX-X	-0.0947
Sant'Angelo	Amatrice RI	IX-X	IX-X	-0.0468
Faizzone	Amatrice RI	IX		-0.0635
Sommati	Amatrice RI	IX	IX	-0.0301
Crognale	Amatrice RI		IX	-0.0790
Accumoli	Accumoli RI	VIII-IX	VIII	-0.1740
Grisciano	Accumoli RI	VIII-IX	VIII-IX	-0.0953
Poggio Casoli	Accumoli RI	VIII-IX	VIII-IX	-0.1651
Cornillo Vecchio	Amatrice RI	VIII-IX	VIII	-0.0689
Cossito	Amatrice RI	VIII-IX	VIII	-0.0564
Retrosi	Amatrice RI	VIII-IX	VIII	-0.0175
Rocchetta	Amatrice RI	VIII-IX	VIII	-0.0959
Arquata del Tronto	Arquata del Tronto AP	VIII-IX	VIII-IX	0.0027
Capodacqua	Arquata del Tronto AP	VIII-IX	VIII-IX	-0.1672
Tufo	Arquata del Tronto AP	VIII-IX	VIII-IX	-0.1557
Fonte del Campo	Accumoli RI	VIII	VII-VIII	-0.2148
Cascello	Amatrice RI	VIII	VII-VIII	-0.0325
Moletano	Amatrice RI	VIII	VIII	-0.0155
Santo Masso	Amatrice RI	VIII		-0.0930
San Giovanni	Accumoli RI	VII-VIII	VIII	-0.0388
Tino	Accumoli RI	VII-VIII	VII-VIII	-0.2043
Collepaggiuica	Amatrice RI	VII-VIII	VII-VIII	-0.0254
Prato	Amatrice RI	VII-VIII	VIII-IX	-0.0478
San Capone	Amatrice RI	VII-VIII		-0.0777
Pretare	Arquata del Tronto AP	VII-VIII	VII-VIII	0.0037
Cossara	Amatrice RI		VII-VIII	-0.0144
Fornisco	Valle Castellana TE		VII-VIII	0.0017
Villanova	Accumoli RI	VII	VI-VII	-0.0865
Capricchia	Amatrice RI	VII	VII	0.0090
Poggio Vitellino	Amatrice RI	VII		-0.0964



San Lorenzo a Pinaco	Amatrice RI	VII		-0.0081
Scai	Amatrice RI	VII	VII	0.0050
Torrta	Amatrice RI	VII		0.0031
Voceto	Amatrice RI	VII	VIII	0.0012
Borgo	Arquata del Tronto AP	VII	VII	0.0015
Faete	Arquata del Tronto AP	VII		0.0088
Piedilama	Arquata del Tronto AP	VII	VII	0.0041
Trisungo	Arquata del Tronto AP	VII	VII	0.0106
Castro	Montegallo AP	VII	VII	0.0042
San Pellegrino	Norcia PG	VII	VII-VIII	-0.0213
Faete	Arquata del Tronto AP		VII	0.0087
Poggio d'Api	Accumoli RI		VII	0.0054
Santa Lucia	Monteale AQ		VII	0.0043
Tallacano	Acquasanta Terme AP		VII	0.0049
Vezzano	Arquata del Tronto AP		VII	-0.0109
Colleposta	Accumoli RI	VI-VII	VI-VII	-0.0075
Falciano	Acquasanta Terme AP	VI-VII		0.0030
Arafranco-Pinaco	Amatrice RI	VI-VII	VII	-0.0039
Castel Trione	Amatrice RI	VI-VII	VII-VIII	0.0082
Cornelle di Sotto	Amatrice RI	VI-VII	VII-VIII	0.0019
Ferrazza	Amatrice RI	VI-VII	VII-VIII	-0.0045
Mosicchio	Amatrice RI	VI-VII	VI-VII	-0.0307
Preta	Amatrice RI	VI-VII	VII	0.0080
San Benedetto	Amatrice RI	VI-VII		-0.0298
San Cipriano	Amatrice RI	VI-VII	VII	-0.0182
Santa Giusta	Amatrice RI	VI-VII	VI	-0.0466
Spelonga	Arquata del Tronto AP	VI-VII	VI-VII	0.0041
Balzo	Montegallo AP	VI-VII	VI-VII	0.0044
Castelluccio	Norcia PG		VI-VII	-0.0607
Collalto	Amatrice RI		VI-VII	-0.0825
Colle	Arquata del Tronto AP		VI-VII	0.0182
Collecetra	Amatrice RI		VI-VII	-0.0112
Configno	Amatrice RI		VI-VII	-0.0100
Nommisci	Amatrice RI		VI-VII	-0.0026
Collegentilese	Amatrice RI	VI	VI-VII	0.0000
Collemoresco	Amatrice RI	VI	VI	-0.0195
Colli	Amatrice RI	VI	VI-VII	-0.0397
Cornillo Nuovo	Amatrice RI	VI	VI	0.0054

Pasciano	Amatrice RI	VI	VI	0.0012
Castelsantangelo sul Nera	Castelsantangelo sul Nera MC	VI	VI	-0.0017
Gualdo	Castelsantangelo sul Nera MC	VI	VII	-0.0047
Noceria	Castelsantangelo sul Nera MC	VI		-0.0010
Norcia	Norcia PG	VI		0.0006
Nottoria	Norcia PG	VI		-0.0072
Ceraso	Valle Castellana TE	VI		0.0020
Morrice	Valle Castellana TE	VI		0.0034
Forcelle	Amatrice RI		VI	0.0031
Pascellata	Valle Castellana TE		VI	0.0021
Quintodecimo	Acquasanta Terme AP		VI	0.0069
Roccasalli	Accumoli RI		VI	0.0012
Varoni	Amatrice RI		VI	0.0049

137 AP: Ascoli Piceno, AQ: L'Aquila, MC: Macerata, PG: Perugia, RI: Rieti, TE: Teramo.

138

139 The sites with maximum InSAR downward displacement were Fonte del Campo, Tino and Illica, all  
140 located in the Municipality of Accumoli, with vertical displacement around -0.20 m.

141 The regression of InSAR vertical displacement and MCS intensity is reported in Figure 4 ( $R^2 =$   
142 0.4631):

143

$$w = -0.0215 I_{MCS} + 0.1133 \quad (w \text{ in m}) \quad (1)$$

144

145 The regression was determined by also considering intensities  $V \leq I_{MCS} < V-VI$ . The higher the  
146 downward displacement, the higher the MCS intensity.

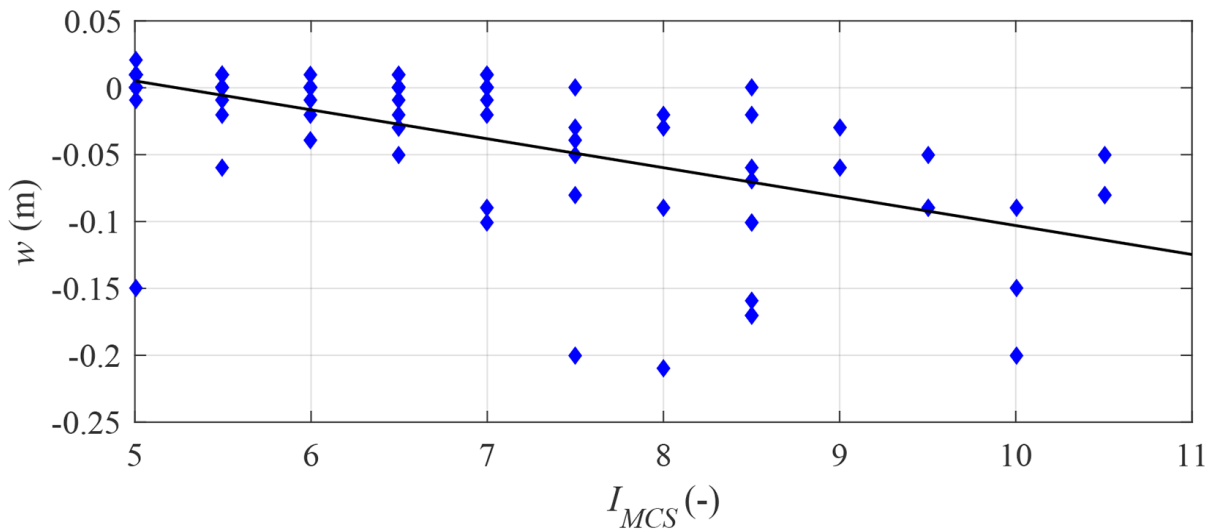
147 It can be observed that downward displacement for the centre of Amatrice is moderate ( $w = -0.04643$   
148 m) but MCS intensity reaches its highest value ( $I_{MCS} = X-XI$ ). Topographic amplification related to  
149 crest morphology and severe damage induced by poor masonry quality (Sorrentino et al. 2018) are  
150 possible explanations. On the other hand, MCS intensities for sites such as Tino and Fonte del Campo,  
151 both in the Municipality of Accumoli, with greatest downward displacement, were moderate, namely  
152 VII-VIII and VIII, respectively. Similarly, downward displacement for the site of Forche Canepine,  
153 in the Municipality of Arquata del Tronto, was high ( $w = -0.14625$  m) with low intensity ( $I_{MCS} = V$ ).  
154 However, this site has few buildings, possibly resulting in a non-robust estimation of intensity.

155 Inverse regression, i.e. calculating the errors along the direction of the  $I_{MCS}$  axis, is:

156

$$I_{MCS} = -21.52 w + 5.714 \quad (w \text{ in m}) \quad (2)$$

157



158

159 Figure 4. Regression between InSAR vertical displacement and MCS intensity (Eq. (1)).

160

161 Similar regressions were derived for EMS intensity (Figure 5), but with larger scatter ( $R^2 = 0.3711$ ):

162

$$w = -0.0214 I_{EMS} + 0.1169 \quad (w \text{ in m}) \quad (3)$$

163

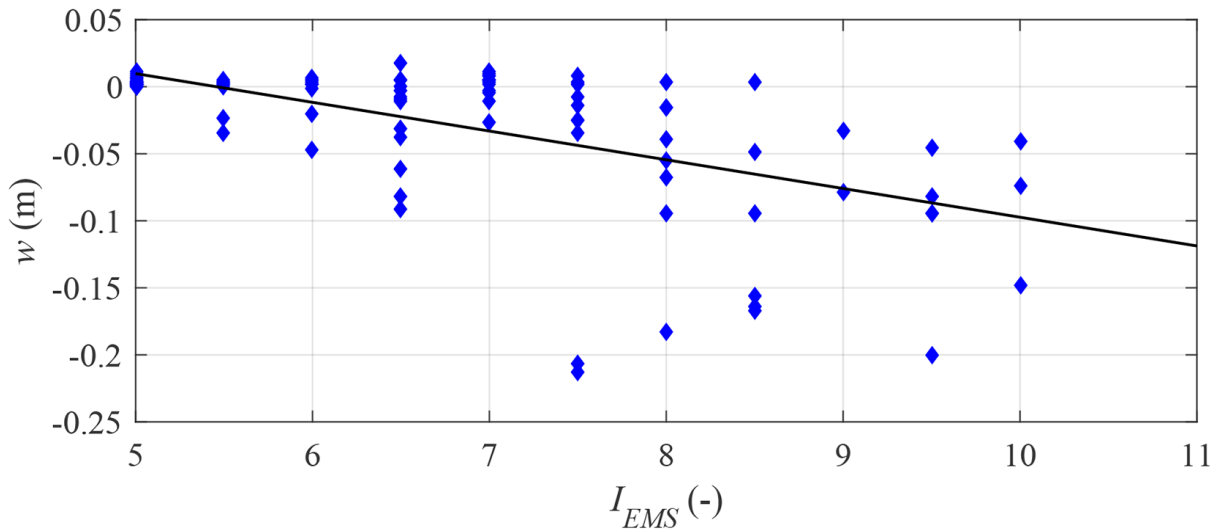
$$I_{EMS} = -17.32 w + 6.134 \quad (w \text{ in m}) \quad (4)$$

164

165 It can be observed that the regressions  $w-I_{MCS}$  and  $w-I_{EMS}$  are almost coincident, whereas the inverse  
166 regressions show significant differences.

167 The inverse regressions (Eqs. 2,4) provide an estimation of MCS and EMS intensity, respectively,  
168 given the InSAR data. However, it must be emphasised that further studies on a statistically  
169 significant set of earthquakes are needed to systematically exploit InSAR data for a quick estimate of  
170 macroseismic intensities in the aftermath of an event.

171



172

173 Figure 5. Regression between InSAR vertical displacement and EMS intensity (Eq. (3)).

174

175 A comparison of MCS and EMS intensities is also interesting. Taking into account only intensity  
 176 values greater than or equal to VI, the  $I_{EMS}$ - $I_{MCS}$  regression is ( $R^2 = 0.8569$ ):

177

$$I_{EMS} = 0.7302 I_{MCS} + 2.173 \quad (5)$$

178

179 resulting in almost equal values for  $I_{MCS} = VIII \cong I_{EMS}$ , whereas for  $I_{MCS} = VI$ ,  $I_{EMS} \cong VI-VII$ , and for  
 180  $I_{MCS} = X$ ,  $I_{EMS} \cong IX-X$ . The inverse regression is:

181

$$I_{MCS} = 1.174 I_{EMS} - 1.472 \quad (6)$$

182

183 It is worth noting that this relationship differs in terms of constant value from the relationship  
 184 proposed by Margottini et al. (1987):  $I_{MCS} = 1.17 I_{EMS} - 0.76$ .

185

186 4 RELATIONSHIPS BETWEEN InSAR VERTICAL DISPLACEMENT AND GROUND  
 187 MOTION INTENSITY MEASURES

188

189 A second group of empirical relationships links InSAR vertical displacement to horizontal ground  
 190 motion intensity measures. These parameters, in turn, can be related to macroseismic intensity by  
 191 means of empirical relationships established in literature. Ground motion intensity measures, as  
 192 opposed to macroseismic intensities, are continuous quantities, are not affected by the intrinsic scatter  
 193 related to structural response, and are not affected by the conventional estimation of damage and

194 vulnerability, and by damage accumulation. On the other hand, the number of records in sites  
195 experiencing downward displacements is rather limited. For the mainshock of August 24, 2016, just  
196 one station (Amatrice) is available and, for the shock of October 30, 2016, five stations, thanks also  
197 to the mobile accelerometric network (Luzi et al. 2017). InSAR data for October 30th cumulate the  
198 effects of the shocks of October 26th and 30th.

199 For the 2016 October 30 event, it can also be observed that, for a given distance from the epicentre,  
200 greater vertical acceleration was recorded in zones with downward displacement than in uplifted  
201 areas.

202 Different ground motion intensity measures are considered, calculated from corrected horizontal  
203 accelerograms. Peak ground acceleration (*PGA*) and peak ground velocity (*PGV*) are selected as  
204 ground motion parameters because they are the most commonly used intensity measures. The ratio  
205 between peak ground acceleration and peak ground velocity (*PGA/PGV*) provides useful information  
206 on frequency content, and is inversely correlated with magnitude, duration, epicentral distance and  
207 the predominant period of the site (Castaldo and Tubaldi 2018). The *PGA/PGV* ratio is expressed in  
208 the form  $PGA / (g PGV)$ , where  $g$  is the gravitational acceleration. Dimensionally, the ratio  $PGA / (g$   
209  $PGV)$  is the inverse of a velocity.

210 Another velocity measure is maximum incremental velocity (*IV*), given by the area below the largest  
211 acceleration pulse (Anderson and Bertero 1987). Peak ground displacement (*PGD*) is not considered  
212 in this study because of the correction procedure consisting of baseline correction, non-causal 2nd  
213 order high-pass and low-pass Butterworth filter, sinusoidal taper and removal of linear  
214 displacement drift (Luzi et al. 2017). The effect of this procedure is a zero final displacement, which  
215 contrasts with physical evidence and InSAR data. A further confirmation of non-zero final  
216 displacement is provided by high frequency GPS records (Wilkinson et al. 2017).

217 Among instrumental intensity measures, Arias, Fajfar and Housner Intensities are taken into  
218 consideration. Arias Intensity ( $I_A$ ) (Arias 1970) is given by:

219

$$I_A = \frac{\pi}{2g} \int_0^{\infty} a_g^2(t) dt \quad (7)$$

220

221 where  $a_g$  is ground acceleration and  $t$  time. Arias Intensity has been proved to represent the sum of  
222 the total energies, per unit weight, stored at the end of the earthquake ground motion in a population  
223 of undamped linear oscillators. Arias Intensity, which dimensionally is a velocity, can be correlated  
224 to damage (Cabañas et al. 1997) but tends to overestimate the intensity of earthquakes with long  
225 duration, high acceleration and broad band frequency content (Uang and Bertero 1988). Arias

226 Intensity has been demonstrated to be an effective predictor of damage to short-period structures  
227 (Stafford et al. 2009).

228 Fajfar Intensity ( $I_F$ ) (Fajfar et al. 1990) is defined as:

229

$$I_F = PGV t_D^{0.25} \quad (8)$$

230

231 where centimetres and seconds are used,  $t_D$  is the Trifunac and Brady strong motion duration (Trifunac  
232 and Brady 1975):

233

$$t_D = t_{0.95} - t_{0.05} \quad (9)$$

234

235 and  $t_{0.05}$  and  $t_{0.95}$  are the time values at which 5% and 95% of the time integral of the history of squared  
236 accelerations are reached, respectively. Fajfar Intensity was formulated to represent earthquake  
237 potential to damage medium-period structures.

238 Housner Intensity ( $I_H$ ) (Housner 1952) is defined as the integral of the elastic pseudo-velocity  
239 spectrum, over the period  $T$  ranging between 0.1 and 2.5 s:

240

$$I_H = \int_{0.1}^{2.5} S_{pv}(T, \xi = 0.05) dT = \frac{1}{2\pi} \int_{0.1}^{2.5} S_{pa}(T, \xi = 0.05) T dT \quad (10)$$

241

242 where  $S_{pv}$  and  $S_{pa}$  are the pseudo-velocity and the pseudo-acceleration, respectively, at undamped  
243 natural period  $T$  and damping ratio  $\xi = 0.05$ . Dimensionally, Housner Intensity is a displacement.  
244 Housner Intensity can be considered as the first moment of the area of  $S_{pa}$  ( $0.1 \text{ s} \leq T \leq 2.5 \text{ s}$ ) about  
245 the  $S_{pa}$  axis. Therefore, it is larger for ground motions with a significant amount of low frequency  
246 content.

247 The analysis has been carried out taking into account the record at Amatrice of August 24, 2016, and  
248 the five records of October 30, 2016, (Luzi et al. 2017) in the zone of downward displacement ( $w \leq$   
249  $-10 \text{ mm}$ ).

250 A set of directions in the horizontal plane, with angular step  $10^\circ$ , has been considered, the aim being  
251 to determine the highest values of horizontal ground motion intensity measures. For each direction,  
252 acceleration time history has been calculated by projecting the NS and EW components, and ground  
253 motion intensity measures have been determined. The maximum values of the parameters are reported  
254 in Table 3.

255

256 Table 3. Ground motion intensity measures in the downward displacement zone.

Station	Date	$M_w$	Soil type	Epical distance (km)	$PGA$ (cm/s <sup>2</sup> )	$PGV$ (cm/s)	$PGA / (g PGV)$ (s/m)	$IV$ (cm/s)	$I_A$ (cm/s)	$I_F$ (cm s <sup>-3/4</sup> )	$I_H$ (cm)	$w$ (m)
AMT	2016/08/24	6.0	B	8.5	851.2	45.18	1.921	79.40	188.2	62.85	121.2	-0.056
CLO	2016/10/30	6.5	A*	7.8	590.5	69.58	0.865	123.31	443.9	117.25	277.4	-0.780
CNE	2016/10/30	6.5	C*	7.7	487.2	41.66	1.192	70.71	197.4	64.92	129.8	-0.198
FCC	2016/10/30	6.5	A*	11.0	938.5	81.76	1.170	139.01	840.2	127.20	233.3	-0.455
IV.T1213	2016/10/30	6.5	A*	12.0	883.5	62.03	1.452	89.48	616.1	99.76	158.7	-0.085
IV.T1214	2016/10/30	6.5	B*	11.4	623.7	56.24	1.130	59.43	399.2	90.70	143.8	-0.426

257

258 The regressions of  $PGA$ ,  $PGV$ ,  $PGA / (g PGV)$ ,  $IV$ ,  $I_A$ ,  $I_F$ ,  $I_H$  vs.  $w$  are:

259

$$PGA = 225.0w + 804.1 \quad (PGA \text{ in cm/s}^2, w \text{ in m}) \quad (R^2 = 0.1124) \quad (11)$$

260

$$PGV = -33.35w + 48.30 \quad (PGV \text{ in cm/s}, w \text{ in m}) \quad (R^2 = 0.3720) \quad (12)$$

261

$$PGA/(g PGV) = 1.108w + 1.658 \quad (PGA/(g PGV) \text{ in s/m}, w \text{ in m}) \quad (R^2 = 0.7138) \quad (13)$$

262

$$IV = -62.38w + 72.77 \quad (IV \text{ in cm/s}, w \text{ in m}) \quad (R^2 = 0.3050) \quad (14)$$

263

$$I_A = -278.4w + 354.7 \quad (I_A \text{ in cm/s}, w \text{ in m}) \quad (R^2 = 0.0937) \quad (15)$$

264

$$I_F = -66.50w + 71.62 \quad (I_F \text{ in cm/s}^{-3/4}, w \text{ in m}) \quad (R^2 = 0.4804) \quad (16)$$

265

$$I_H = -197.4w + 111.6 \quad (I_H \text{ in cm}, w \text{ in m}) \quad (R^2 = 0.7406) \quad (17)$$

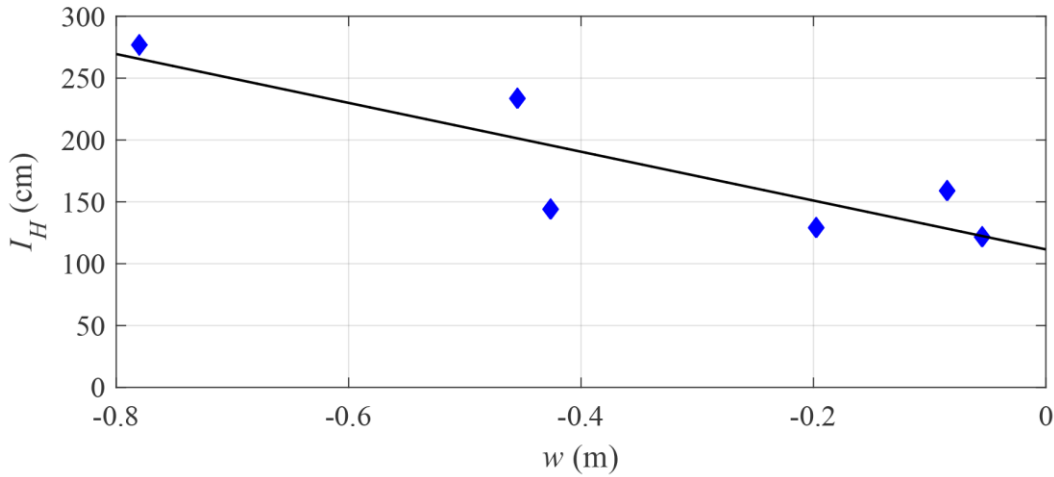
266

267 The regressions only hold in the downward displacement zone ( $w \leq -10$  mm) and cannot be extended  
 268 to the upward or negligible vertical displacement zones. They provide a preliminary indication of the  
 269 relationship between InSAR vertical displacement and ground motion intensity measures.

270 It can be observed that the correlation of InSAR displacement with the parameters related to high  
 271 frequency content ( $PGA$ ,  $I_A$ ) is limited. It can also be observed that  $PGA$  is positively correlated with  
 272  $w$ , and this further confirms the limited reliability of the  $PGA$ - $w$  regression.

273 InSAR vertical displacements are moderately correlated with the parameters related to intermediate  
 274 frequency content ( $PGV$ ,  $IV$ ,  $I_F$ ), and well correlated with the ratio  $PGA / (g PGV)$  and Housner  
 275 Intensity  $I_H$ , (Figure 6), the latter being mainly related to low frequency content. Regarding Housner  
 276 Intensity, studies in literature (Decanini et al. 2002, Masi et al. 2011, Marotta et al. 2017)

277 demonstrated that it can be a valid alternative to other seismic peak parameters, and in (Chiauzzi et  
278 al. 2012) a relationship between Housner Intensity and EMS intensity was developed.  
279



280  
281 Figure 6. Regression between Housner Intensity and InSAR vertical displacement.

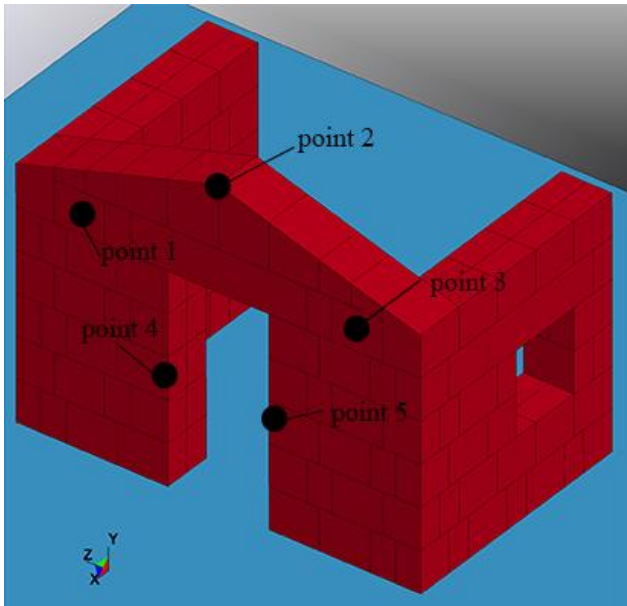
282  
283

## 284 5 EFFECT OF THE VERTICAL COMPONENT ON THE SEISMIC RESPONSE OF A 285 MASONRY TEST STRUCTURE

286  
287 The effect of the vertical component on the seismic response of a masonry structure has been  
288 investigated by means of numerical simulations of the response of a test structure in natural stone  
289 shown in Figure 7. The façade is 4.15 m in width and reaches a maximum height of 3.00 m; the lateral  
290 walls are 2.50 m wide and 2.45 m high. Thickness for both façade and lateral walls is 0.50 m. This  
291 structure had been previously tested physically (Candeias et al. 2017) and numerically (Mendes et al.  
292 2017, AlShawa et al. 2017).

293





294

295 Figure 7. Masonry test structure.

296

297 The test structure was modelled by means of a combined finite-discrete element strategy implemented  
298 in LS-DYNA (Hallquist 2006). The model has discrete components, i.e. block elements and contact  
299 interfaces. The three-dimensional model was able to capture the experimentally observed  
300 mechanisms, without any *a priori* assumption about the mechanism itself.

301 The modelling approach for the block elements allows for cracking, separation, and re-contact along  
302 predefined contact surfaces. In order to avoid significant bias in the failure mechanism, the interfaces  
303 must be plentiful, and at the same time must be limited in number to contain modelling and  
304 computational burden. Consequently, a block element represents several masonry units and mortar  
305 joints. Linear-elastic behaviour is assumed for the material. Young's modulus,  $E$ , and density,  $\rho$ , of  
306 the units have been set at:  $E = 2077$  MPa,  $\rho = 2360$  kg/m<sup>3</sup>. Poisson's ratio has been set at 0.2. Damping  
307 ratio has been assumed to be 0.05. The standard unit in the model is  $500 \times 250 \times 400$  mm<sup>3</sup>. The blocks  
308 are parallelepipeds, more regular than the real ones. The number of blocks and of FE total 169 and  
309 7716, respectively (AlShawa et al. 2017).

310 Regarding contact interfaces, automatic detection of contact is implemented. Explicit time integration  
311 is adopted, providing more stable results when contact interfaces are present (Burnett et al. 2007,  
312 Jäger et al. 2009, Srewil 2008). Interfaces transmit both compression and tension, with optional  
313 failure criteria for the latter (Bala 2007). Tension is resisted by a linear contact spring until failure.  
314 Spring failure criterion is based on normal force alone, that is, the spring is removed when:

315

$$\sigma_n \geq NFLS \tag{18}$$

316

317 where  $\sigma_n$  is the normal stress on the contact surface and *NFLS* the normal (tensile) failure limit stress.

318 Post-failure interaction takes place according to classical compression-friction contact. Initiation of  
319 sliding is controlled by shear failure limit stress in the absence of normal stress, *SFLS*, representing  
320 the interface cohesion, and by friction, governed by Coulomb's law. The analyses assume *SFLS* =  
321 *NFLS*, based on a pure shear condition (Calderini et al. 2010), where *SFLS* = *NFLS* = 0.010, 0.150  
322 MPa (AlShawa et al. 2017), the lower value accounting for smooth units and decayed mortar, as  
323 commonly found in many buildings damaged/collapsed following the Amatrice-Norcia earthquake.  
324 Once shear capacity has been exceeded, sliding is controlled by static and dynamic friction  
325 coefficients. For both static and dynamic friction coefficients, values  $\mu = 0.4, 0.6, 0.8$  have been  
326 adopted, based on experimental values reported in Burnett et al. (2007), Maheri et al. (2011),  
327 Liberatore et al. (2016), the lowest value representing the masonry characteristics of many buildings  
328 involved in the Amatrice-Norcia earthquake. The dynamic friction coefficient has been assumed to  
329 be equal to the static coefficient because of lack of experimental data on the types of masonry  
330 investigated.

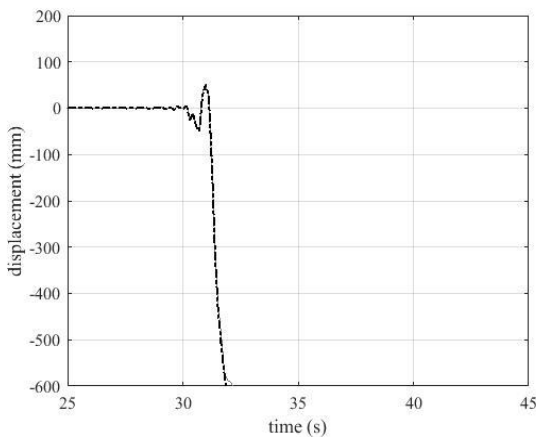
331 The out-of-plane displacement time histories of point 2 (Figure 7) under CLO record of 30th October  
332 2016 (see Table 3), which is associated with the maximum downward displacement ( $w = -0.780$  m),  
333 are shown in Figures 8 and 9. Either the NS or EW component was applied along the direction  
334 orthogonal to the façade. The vertical component is assumed to be either acting or not.

335 It can be noticed that the vertical component produces displacement shifts in some time histories for  
336 *SFLS* = *NFLS* = 0.010 MPa ( $\mu = 0.4$ , EW;  $\mu = 0.6$ , NS;  $\mu = 0.8$ , EW), indicating an increase in sliding  
337 at unit-mortar interfaces (Figure 8).

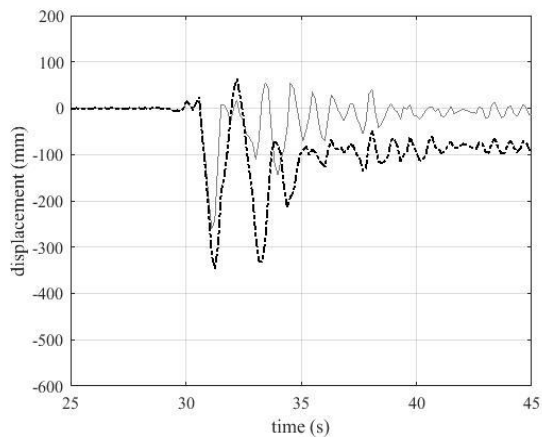
338 The minimum (negative, inwards) and maximum (positive, outwards) values of displacement at  
339 points 1-5 and base shear are reported in Figures 10-13. Displacements in absolute value greater than  
340 200 mm are associated with large out-of-plumb, corresponding to failure – or near failure – condition  
341 of the wall. It can be noticed that a number of failures occur for *SFLS* = *NFLS* = 0.010 MPa, consistent  
342 with the post-earthquake reconnaissance of damage. No failures occur for *SFLS* = *NFLS* = 0.150  
343 MPa. Increases in base shear can be observed for *SFLS* = *NFLS* = 0.150 MPa compared to *SFLS* =  
344 *NFLS* = 0.150 MPa.

345 It can be noticed that response under the NS component is generally higher than that under the EW  
346 component, in terms of both displacements and base shear, consistent with the corresponding  
347 response spectra (Luzi et al. 2017).

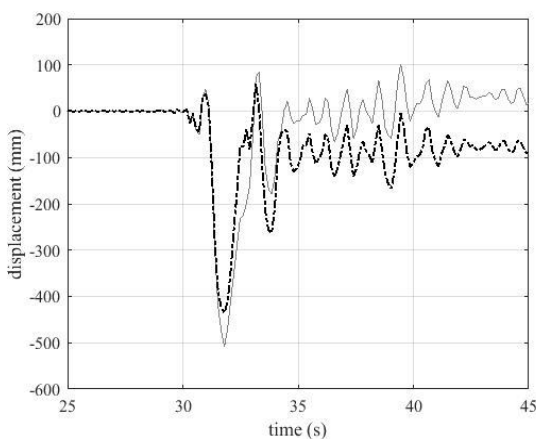
348 The vertical component induces a set of additional failures for  $SFLS = NFLS = 0.010$  MPa, namely:  
349  $\mu = 0.4$ , NS (points 4, 5), EW (point 3);  $\mu = 0.8$ , EW (point 3), and base shear shows a mean increment  
350 of 9%. On the other hand, for  $SFLS = NFLS = 0.150$  MPa, the effect of the vertical component is  
351 generally limited, with a mean relative decrement of 3% in displacements and an increment of 1% in  
352 base shear. These limited effects can be explained on the basis of the frequency content of the vertical  
353 ground acceleration, greater in the high frequency range, compared to displacement along horizontal  
354 directions. During a single sliding along a mortar joint, several cycles of normal (vertical) stress  
355 occur, along with a corresponding increase/decrease of friction force, resulting in a non-systematic  
356 effect on the total sliding.  
357



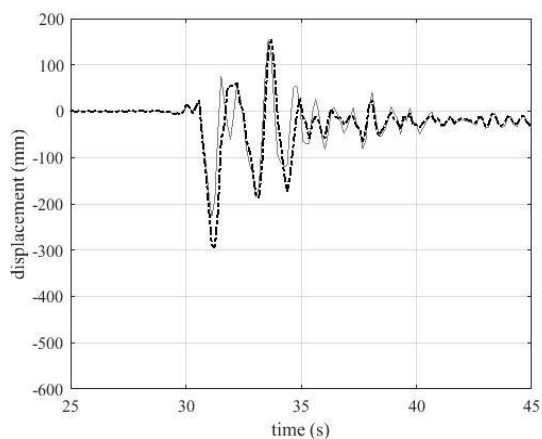
358  
359  
360  
a) NS,  $\mu = 0.4$



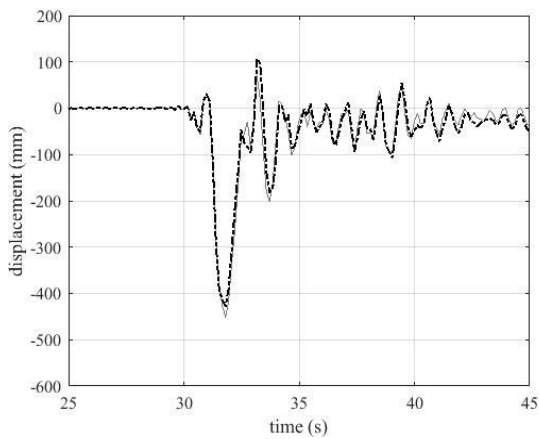
b) EW,  $\mu = 0.4$



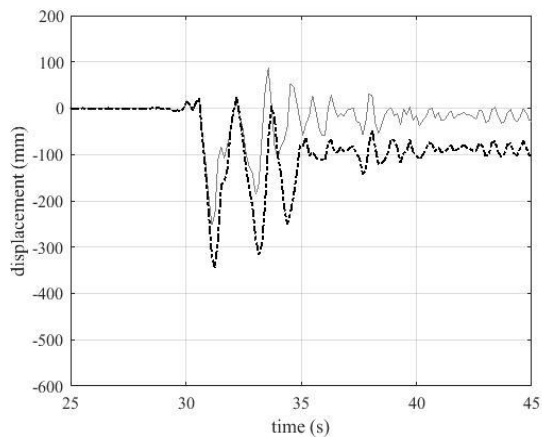
361  
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363  
c) NS,  $\mu = 0.6$



d) EW,  $\mu = 0.6$



e) NS,  $\mu = 0.8$



f) EW,  $\mu = 0.8$

364

365

366

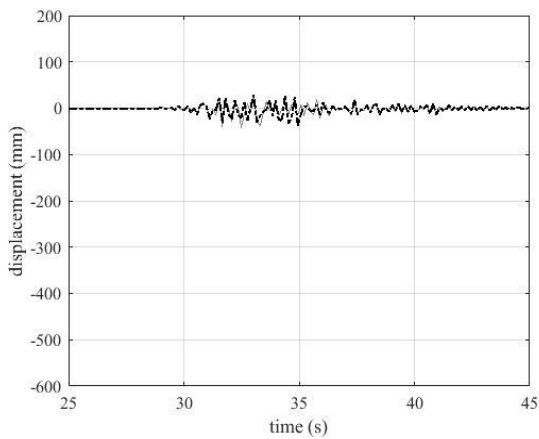
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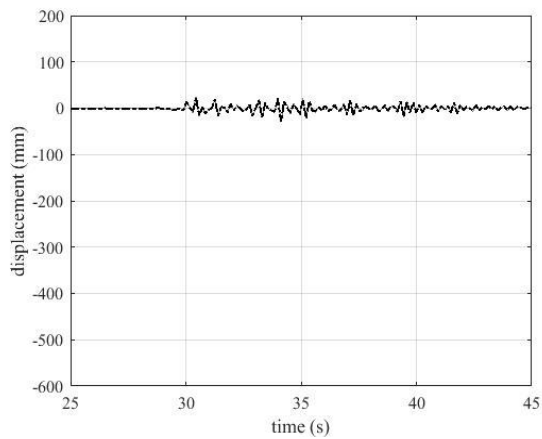
369

370

Figure 8. Out-of-plane displacement time histories of point 2 of the test structure subjected to the CLO record of 30th October 2016,  $SFLS = NFLS = 0.010$  MPa (grey: without the vertical component; black: with the vertical component).



a) NS,  $\mu = 0.4$

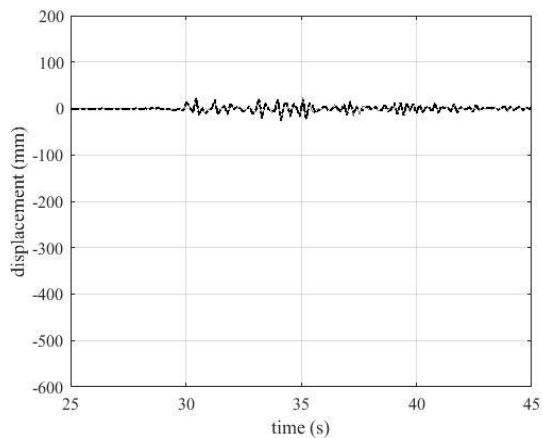
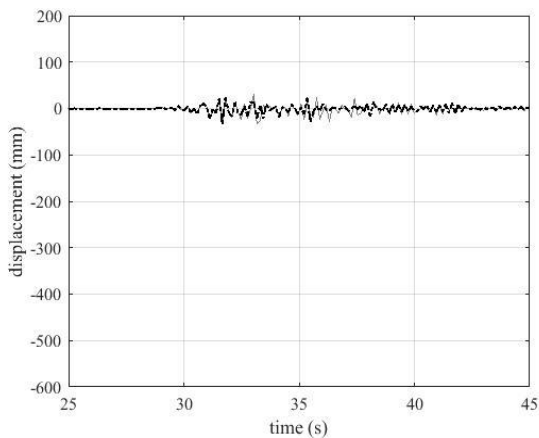


b) EW,  $\mu = 0.4$

371

372

373

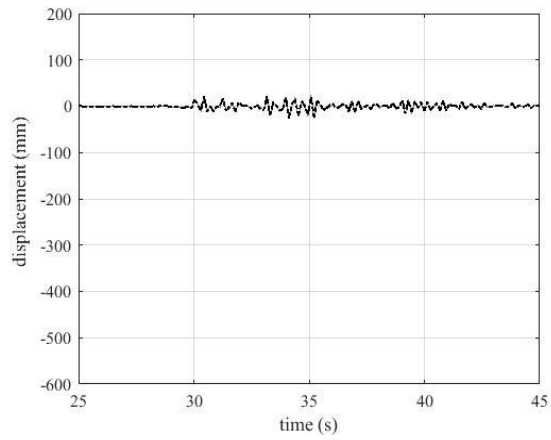
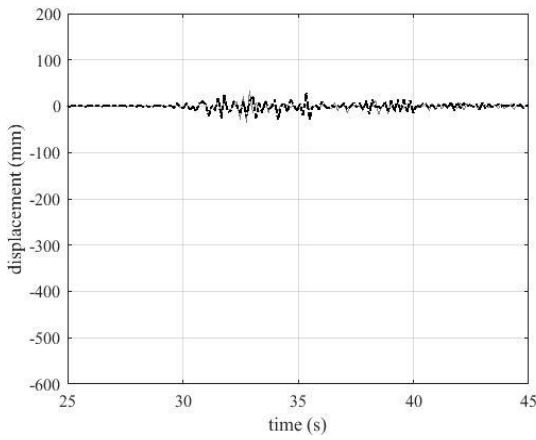


374

375  
376

c) NS,  $\mu = 0.6$

d) EW,  $\mu = 0.6$

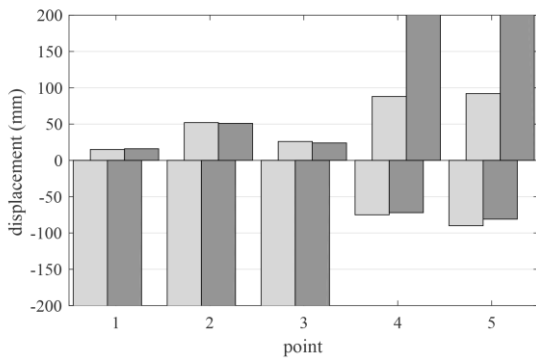


377  
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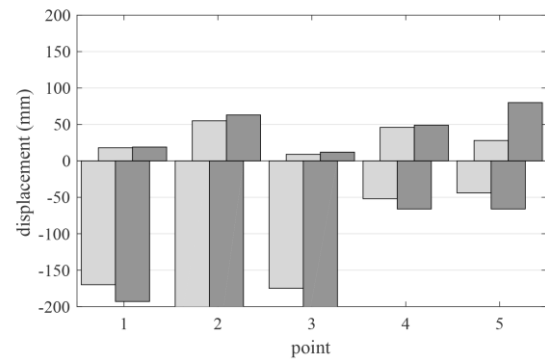
e) NS,  $\mu = 0.8$

f) EW,  $\mu = 0.8$

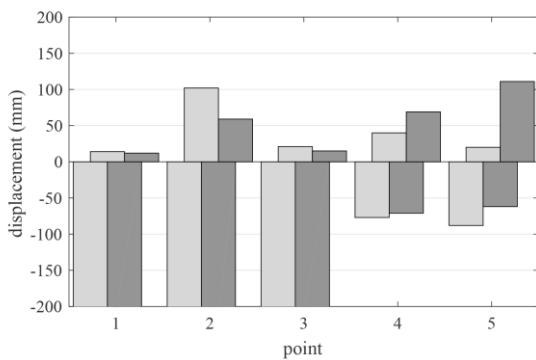
380 Figure 9. Out-of-plane displacement time histories of point 2 of the test structure subjected to the  
381 CLO record of 30th October 2016,  $SFLS = NFLS = 0.150$  MPa (grey: without the vertical  
382 component; black: with the vertical component).  
383



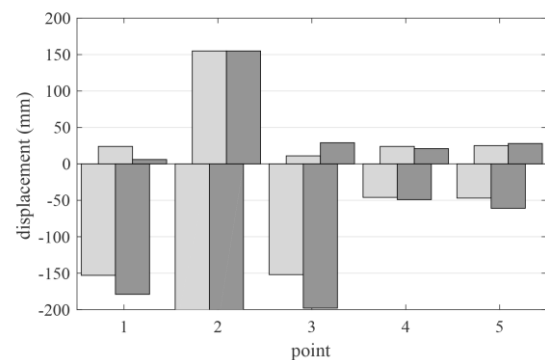
a) NS,  $\mu = 0.4$



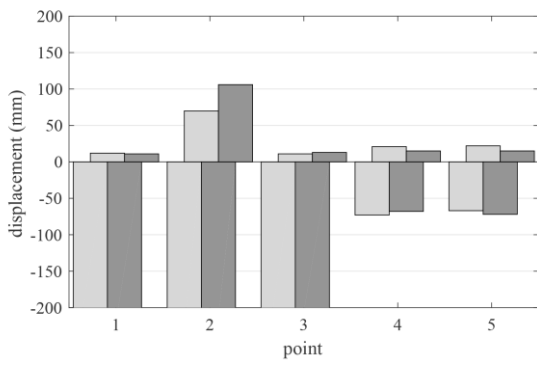
b) EW,  $\mu = 0.4$



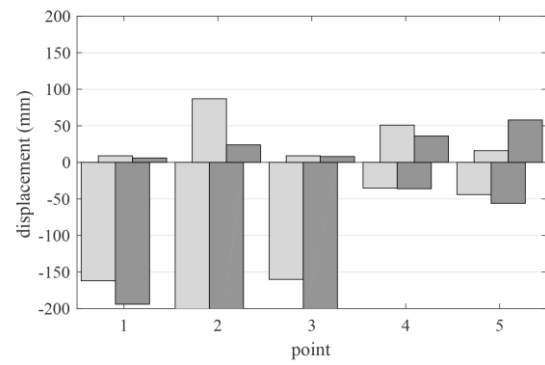
c) NS,  $\mu = 0.6$



d) EW,  $\mu = 0.6$

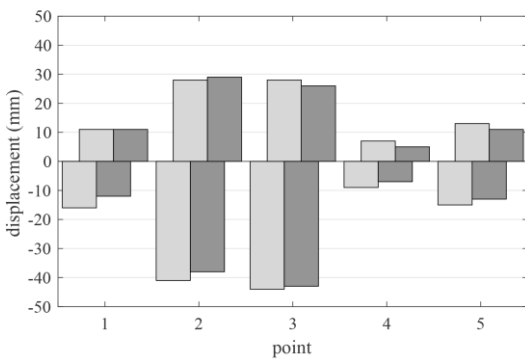


e) NS,  $\mu = 0.8$

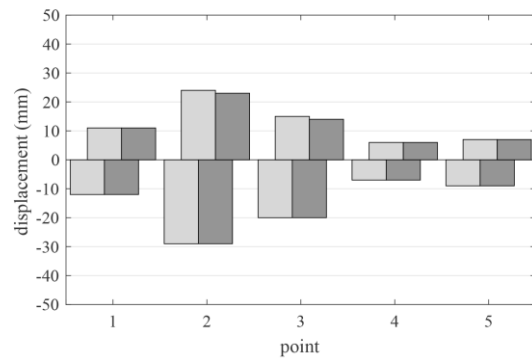


f) EW,  $\mu = 0.8$

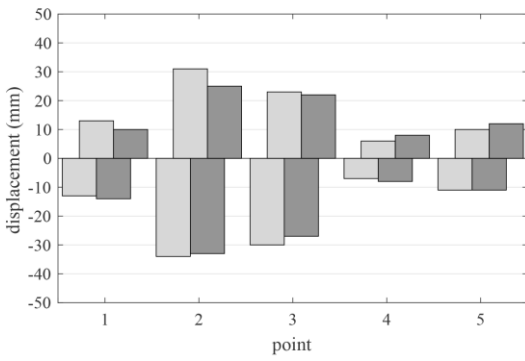
384 Figure 10. Minimum/maximum displacements of the test structure subjected to the CLO record of  
 385 30th October 2016,  $SFLS = NFLS = 0.010$  MPa (light grey: without the vertical component; dark  
 386 grey: with the vertical component).  
 387



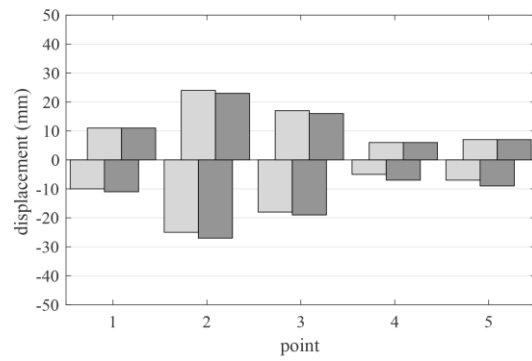
a) NS,  $\mu = 0.4$



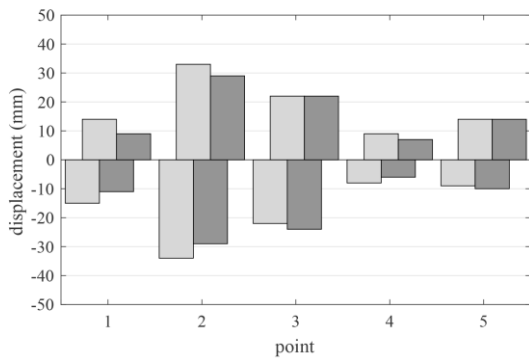
b) EW,  $\mu = 0.4$



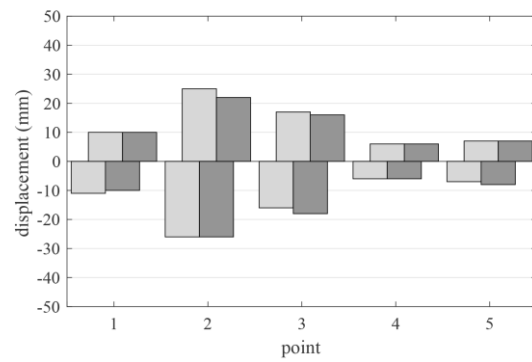
c) NS,  $\mu = 0.6$



d) EW,  $\mu = 0.6$



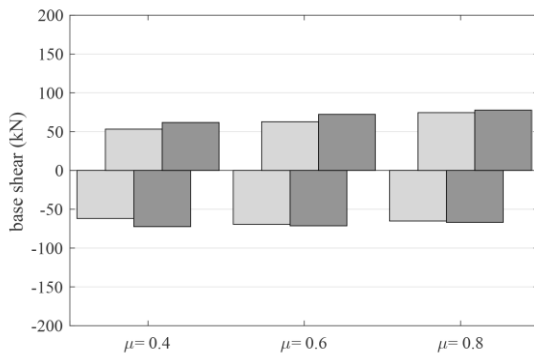
e) NS,  $\mu = 0.8$



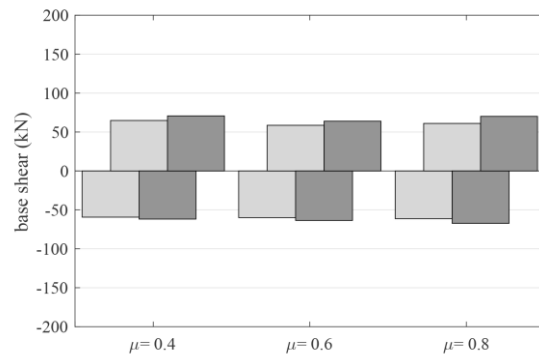
f) EW,  $\mu = 0.8$

388 Figure 11. Minimum/maximum displacements of the test structure subjected to the CLO record of  
 389 30th October 2016,  $SFLS = NFLS = 0.150$  MPa (light grey: without the vertical component; dark  
 390 grey: with the vertical component).

391



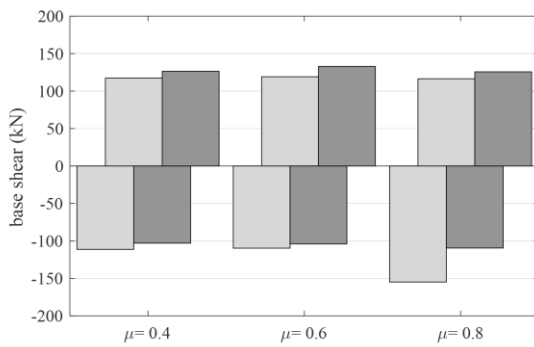
a) NS



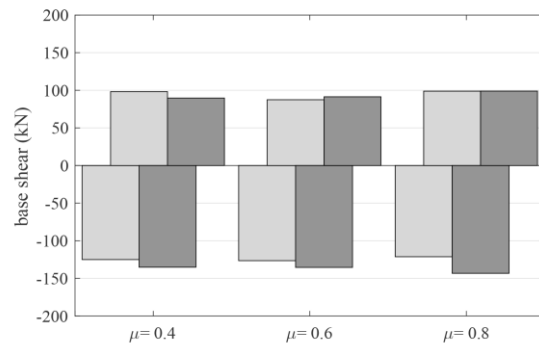
b) EW

392 Figure 12. Minimum/maximum base shear of the test structure subjected to the CLO record of 30th  
 393 October 2016,  $SFLS = NFLS = 0.010$  MPa (light grey: without the vertical component; dark grey:  
 394 with the vertical component).

395



a) NS



b) EW

396 Figure 13. Minimum/maximum base shear of the test structure subjected to the CLO record of 30th  
397 October 2016,  $SFLS = NFLS = 0.150$  MPa (light grey: without the vertical component; dark grey:  
398 with the vertical component).

399

400

## 401 6 CONCLUSIONS

402

403 This paper presents an analysis of InSAR data for the 2016 Amatrice-Norcia earthquake. The research  
404 stemmed from the observation that the most severe damage is localized in the area that underwent  
405 subsidence, whereas zones closer to the epicentre, but coseismically uplifted, experienced less  
406 destruction. This observation is confirmed by statistical analyses, showing good correlation between  
407 InSAR downward displacement and macroseismic intensity.

408 The effect of vertical acceleration, and the ensuing variation of friction forces between units and  
409 mortar in a masonry building, is investigated by means of numerical simulations on a test structure  
410 with cohesive-frictional unit-mortar interface. The test structure was subjected to the horizontal  
411 component only, and to the horizontal component combined with the vertical component. The vertical  
412 component induces more extensive failures in structures with small cohesion, whereas it has limited  
413 effects in structures with medium cohesion. This behaviour can be ascribed to the greater high-  
414 frequency content of vertical ground acceleration, compared to horizontal displacement. During a  
415 single sliding along a mortar joint, several cycles of normal (vertical) stress occur at the interface,  
416 with increasing/decreasing friction force, resulting in non-systematic effects on the total sliding.  
417 However, during the time intervals when friction is lower, larger sliding occurs, with greater masonry  
418 damage.

419 The correlation between InSAR displacement and damage can be explained on the basis of ground  
420 motion intensity measures. In fact, statistical analyses showed good correlation between InSAR  
421 vertical displacement on the one hand, and the ratio  $PGA/PGV$  and Housner Intensity on the other.  
422 Housner Intensity, in particular, has been proved by literature studies to be well correlated to  
423 macroseismic intensity and structural damage.

424 Future developments of this study will investigate the relationships of macroseismic intensity and  
425 ground motion intensity measures vs. InSAR vertical displacement for earthquakes with different  
426 characteristics, in particular reverse focal mechanism and upward displacement in the epicentral zone.

427

428



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430

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432 providing the location of macroseismic data for the 2016 seismic sequence.

433

434

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